

EVALUATION OF FIELD CPTU DISSIPATION DATA IN OVERCONSOLIDATED FINE-GRAINED SOILS

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BACKGROUND

Cone penetration testing with pore pressure measurement (CPTU or piezocone testing) has become a popular investigation technique in geotechnical site investigation practice. The near-continuous data obtained during penetration (approx. every 5 cm) provide information related to the following soil response parameters: tip resistance (q_c), sleeve friction (f_s), and penetration pore pressure (u). In addition, inclination and temperature may also be recorded depending on the type of piezocone being used (Campanella and Robertson, 1988). The penetration of the piezocone can be halted at any depth and the variation with time of the measured parameters can be monitored. Of the above three quantities (q_c , f_s , u), it is usually the variation of the pore pressure that is of interest, as the results can be interpreted to provide estimates of the *in situ* horizontal coefficient of consolidation, c_h (Torstensson 1977). Rather than the total pore pressure, it is the change in the excess pore pressure (Δu) with time that is required for the evaluation of c_h , where Δu is defined as:

$$\Delta u = u_i - u_o \tag{1}$$

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and

u_i is the measured pore pressure at the depth of interest

u_o is the equilibrium *in situ* pore pressure at the depth of interest.

Interpretation of dissipation records is generally based on a normalized excess pore pressure ratio, U , defined as:

$$U = \Delta u(t) / \Delta u_i = [u(t) - u_o] / (u_i - u_o) \quad (2)$$

where:

$\Delta u(t)$ is the excess pore pressure at any time t after penetration is stopped

Δu_i is the initial excess pore pressure at $t=0$, i.e. on stopping penetration

$u(t)$ is the total pore pressure at any time t .

Hence, for standard dissipation records where the excess pore pressure shows a monotonic decrease with time, U varies between unity (at $t=0$) and zero when 100% dissipation of the excess pore pressure has occurred.

With the development of piezocone equipment, it is now possible to measure penetration pore pressures at one or more locations on the cone. Three specific locations will be discussed in this paper and designated according to the scheme illustrated in Fig.1. Hence, the excess pore pressure can also be subscripted according to where the measurements are obtained (Sully et al. 1988):

$$\Delta u_{1,2,3} = u_{1,2,3} - u_o \quad (3)$$

Because of the small distance between each of the three measurement locations (0.1m max. between u_1 and u_3), the magnitude of u_0 can be taken to be equal for each of the three positions.

Once the required dissipation data have been obtained during the penetration testing, the excess pore pressure variation with time can be plotted if the equilibrium pore pressure is known (or measured). In low permeability soils, it is usual practice to continue taking dissipation measurements until at least half the initial excess pore pressure has dissipated ($U = 0.5$).

Interpretation of the dissipation results can be achieved using either of the two main analytical approaches: cavity expansion theory or the strain path approach. Interpretation by dislocation methods has also been proposed by Elsworth (1993). Comparisons of the available solutions and results from field studies suggest that the cavity expansion method of Torstensson (1977) and the strain path approaches of Levadoux (1980) and Teh (1987) all provide similar predictions of consolidation parameters from CPTU dissipation data (Gillespie 1981, Kabir and Lutenecker 1990, Robertson et al. 1991). Robertson et al. (1991) have shown that these methods, although developed for normally consolidated soils, can be equally applied to overconsolidated soils. Furthermore, comparisons of field and laboratory data indicate that the trends in the measured (laboratory) and predicted (CPTU) data are consistent provided the microfabric and nature of the soils being tested are taken into consideration (Danziger 1990, Robertson et al. 1991). Limited published data are available to verify the more recent dislocation method.

However, the relevance of any of the above solutions depends on many factors, the most important of which relates to how well the initial pore pressure distribution around the cone compares with the theoretical idealization employed by each of the models. The initial distribution around the probe may be such that

the applicability of these methods may be questioned or restricted only to normally consolidated soils.

A typical set of excess pore pressure dissipations in soft normally consolidated clay for the three above-mentioned filter locations are presented in Fig. 2(a). The corresponding normalized dissipation curves are shown in Fig. 2(b). All three pore pressure dissipation curves show a monotonic decrease in the excess pore pressure with time and essentially agree with the theoretical models of the dissipation curve. Under these conditions, the data can be interpreted according to any of the available theories to estimate the *in situ* consolidation parameter, c_h , which primarily governs the rate of dissipation for CPTU tests (Baligh and Levadoux 1980). The rate of dissipation is highest on the face of the cone and reduces with distance behind the tip - these effects are considered in the analysis by varying the time factor, T , according to location of the pore pressure element and by using the radius of the probe at the location of the pore pressure measuring sensor in the numerical calculations.

Under certain circumstances, the pore pressures measured behind the tip do not decrease immediately on stopping penetration; rather, they show an initial increase over a definite period of time before finally beginning to dissipate. Where the pore pressure measurement system is completely saturated, dissipation records of this type are characteristic to filter locations located behind the cone tip (u_2 and u_3) for penetration in overconsolidated soils. The interpretation of these dissipation records to obtain predictions of c_h is the subject of this paper.

PORE PRESSURE DISSIPATION IN OVERCONSOLIDATED SOIL

A typical example of pore pressure dissipation in a lightly overconsolidated fine grained soil ($OCR=4$) is illustrated in Fig. 3 using results obtained at Strong Pit in the Lower Mainland of British Columbia (Campanella et al. 1988). Similar types of dissipation record in OC soils for locations behind the tip have been reported by Tumay et al. (1981), Davidson (1985), BRE/NGI (1985), Kabir and Lutenegger (1987), Gillespie et al. (1987), Coop (1987), Lutenegger and Kabir (1988), Lunne et al. (1986), Gomez and Escalante (1987), Campanella et al. (1988) and Bond and Jardine (1991).

The initial rise in pore pressures measured at locations behind the tip in overconsolidated soils may be explained by one or more of the following:

- a) poor saturation and/or poor response of the measurement system, such that a time lag in response to pore pressure changes occurs,
- b) redistribution of pore pressure around the tip due to the large gradients that are generated in OC soils.

(The Mandel-Cryer effect is not considered to be of major importance for locations behind the tip for the overconsolidated soils examined here.)

A further possibility for the increase in pore pressure has been suggested by Coop and Wroth (1989) as a result of the maximum penetration pore pressure being located at some point away from the shaft of the piezocone. These points will be considered briefly below.

Poor saturation and/or poor response of measurement system

If the rise in pore pressure on halting penetration were due solely to saturation and/or measurement problems, then the effects should be equally frequent in normally consolidated as well as overconsolidated soils. This is not the case.

Furthermore, data of this type have been reported by some of the main research centres around the world where OC soils have been studied and experimental techniques are well proven. It would thus appear that the anomalous pore pressure rise is not due solely to poor field technique. However, it must also be borne in mind that in heavily overconsolidated soils pore pressures behind the cone tip may become negative and in some instances give rise to cavitation of the measuring system (Powell et al. 1988). If cavitation occurs the measurement system may become de-saturated and sluggish response will result, giving rise to curves somewhat similar to those shown in Fig. 3.

Redistribution of pore pressure

Due to the large gradient of pore pressures around the tip in overconsolidated and stiff fine grained soils (Robertson et al. 1986, Davidson 1985), drainage from the tip (high pore pressure) to the zone behind the tip (lower pore pressure) occurs, the rate of which is determined primarily by the soil permeability and the magnitude of the gradient. The effect of the flow around the tip on the pore pressures measured on the shaft also varies according to the soil stiffness and strength, parameters which also determine the pore pressure gradient itself. The authors consider this to be the principal reason why pore pressures measured at locations behind the tip in OC and stiff fine grained soils show an initial increase when penetration is stopped followed by dissipation of excess pore pressures. It is also interesting to note that a comparison of the u_2 and u_3 pore pressures in Fig. 3 indicates that the u_2 value reaches a peak faster than the u_3 measurement. This is logical if the driving force for the initial increase is the pore pressure gradient, since the gradient between the u_1 and u_2 locations is much higher than that between the u_2 and u_3 locations (Robertson et al. 1986, Sully et al. 1988, Whittle et al. 1991).

Maximum pore pressure located away from shaft

Coop and Wroth (1989) have suggested that the maximum penetration pore pressure in overconsolidated soils is located at some distance away from the shaft of the penetrometer. If this were the case, soil strength and stiffness would control not only the magnitude of the pore pressure behind the tip but also the distance of the maximum value from the shaft. Hence, in NC soils the maximum pore pressure would be close to or on the shaft and pore pressure decrease would occur on stopping penetration. As the soil became more overconsolidated, the location of the maximum pore pressure would move away from the shaft and progressively longer time delays would occur before the pore pressure reached its peak value after stopping penetration. This idea would imply however, that at all locations on the shaft, irrespective of the distance behind the tip, the measured pore pressures would all attain peak values at the same time. This is not the case, as shown in Fig. 3, and in other OC dissipation data referenced above. General trends in published data also do not confirm the hypothesis of Coop and Wroth (1989).

TYPICAL DISSIPATION RECORDS IN OVERCONSOLIDATED SOILS

Considering the data presented in Fig. 3 it is apparent that only the u_1 pore pressure shows initial dissipation on halting penetration. However, the unloading of the tip resistance causes a sudden decrease in the u_1 pore pressure. This sudden decrease modifies the dissipation record such that normalization with the initial maximum penetration pore pressure $(u_1)_i$ of 247.3 mH_2O^5 gives rise to a non-standard dissipation record compared to that suggested by the available

⁵ 1 mH_2O = 10 kPa = 1 t/m^2

theories. For the behind the tip locations, the pore pressure initially increases when penetration is stopped, before finally decreasing and arriving at the in situ equilibrium value.

In conclusion, it appears that interpretation of pore pressure dissipation records in OC soils is complicated by unloading effects and redistribution at all three filter locations considered here.

Commonly, the dissipation results are presented in terms of normalized curves, whereby the normalized pore pressure, U , at any time t , is given by Eq. (2). The normalized dissipation curves for the records on Fig. 3 are shown on Fig. 4. None of the curves follow the theoretical dissipation trends suggested by the available theories for normally consolidated soils and hence cannot be evaluated, as is, to provide information on the coefficient of consolidation of the soil. The departure of the field curves from the theoretical framework is considered to be a result of the mean normal stress reduction that occurs as the soil passes around the cone tip and along the shaft.

Characteristic Dissipation Types

In normally consolidated soils, the pore pressure dissipation curve for any filter location (Fig. 2) can be considered as a Type I response. A Type I response implies a monotonic decrease of the initial penetration excess pore pressure. In overconsolidated soils, several different responses may be obtained depending on soil characteristics and filter location. For the filter located on the cone tip (u_1), the unloading type dissipation associated with overconsolidated soils can be classified as a Type II response (Fig. 5a). Type II response is similar to Type I once the pore pressure reduction due to unloading has occurred.

Type III response is assigned to the behind the tip locations (u_2, u_3) where the excess pore pressure is greater than hydrostatic ($u(t) > u_0$), but increases on

stopping penetration before dissipating (Fig. 5b). All of the above responses have been discussed earlier.

In moderately to heavily OC soils, the pore pressures measured at the location immediately behind the tip may be less than hydrostatic, or even below zero. In this case, on halting penetration the pore pressure increases to finally arrive at the in situ equilibrium value. Two types of dissipation curve may result depending on the soil characteristics:

- the measured pore pressure may increase over and above the in situ equilibrium value if the rate of pore pressure redistribution is higher than the rate of dissipation. After reaching some peak value, the pore pressure then decreases until the equilibrium value is reached (Fig. 5c, Type IV). The Type IV curve is similar to the Type III response, the difference being the degree of pore pressure change or unloading that occurs.

- if the rate of dissipation is faster than the rate of redistribution, the pore pressure dissipation does not overshoot but directly arrives at the equilibrium value (Fig. 5c, Type V).

The standard approach for interpreting dissipation records in normally consolidated soils cannot be applied to the OC responses (with the exception of an inverted Type V), since dissipation does not follow the theoretical response, that is, a monotonic reduction with time. The theoretical framework for evaluating dissipation in soils showing responses similar to those defined by Types II to IV is not available at present. However, the anomalous curves can be corrected to permit interpretation using the available theories. In fact, Elsworth (1993) suggests curve correction prior to applying the dislocation method for curves where pore pressure increase on stopping penetration is recorded. The suggested correction technique for interpreting these records is outlined below.

The Type V response can be considered as an inverted dissipation and treated in the standard way (as dissipation of a negative excess). This type of response may occur in soils with an OCR of 4 or larger, although the type of response also depends on soil type and structure. For example, in heavily OC Taranto Clay (OCR=20-40) the pore pressures behind the tip are positive, possibly due to the cementation present (Battaglio et al. 1986), whereas in fissured London Clay (OCR=25-50) negative pore pressures are recorded (Powell and Uglow 1988).

CORRECTION TO DISSIPATION CURVES FOR REDISTRIBUTION EFFECTS

Two data manipulation approaches can be utilized to correct the Type II to IV dissipation curves, so that the available dissipation theories can be used to estimate values of the coefficient of consolidation. One approach is based on a log-time plot while the other is based on a square-root time representation, both similar to the routine methods presently employed for laboratory consolidation data. Either approach can be used separately or combined to provide a check on the results obtained. The application of the methods are presented using the results in Fig. 4 as an example. The step-by-step correction procedure for both methods is outlined in Appendix 1.

Log-time plot correction

The data in Fig. 4 have to be corrected according to the location of the pore pressure measurement, i.e. either on the cone tip or behind the cone tip, since the unloading and/or redistribution that occurs affects the three sets of pore pressures in different ways.

On the tip, a sudden decrease in pore pressure occurs on halting penetration. In Fig. 4 the normalized pore pressure 5 seconds after dissipation begins is already reduced by 25% due to the reduction in the bearing stress acting on the face of the cone. For this location, the initial maximum pore pressure used for normalizing the dissipation record is taken as the peak value once the initial unload has occurred (u_c), i.e. for this case the maximum value corresponds to the 5 sec. measurement and this time (t_c) is taken as the new zero time point (5 sec. are subtracted from the time register throughout the record). The maximum pore pressure for the dissipation record is taken as the peak value which occurs, for this particular record, at 5 seconds.

For the behind tip locations, the maximum pore pressure is taken as the peak value that occurs during the post-penetration increase and the time at which this peak occurs is taken as the new zero time of the dissipation record and all other times adjusted accordingly.

The data from Fig. 4, corrected in this way, are replotted in Fig. 6 to show the new form of the normalized dissipation plot, adjusted to account for unloading and redistribution effects. The method of correcting the data is considered to be theoretically acceptable since it adjusts the dissipation file so that a monotonic reduction of a maximum pore pressure to the in situ equilibrium occurs thereby removing the anomalous effects discussed above.

Root-time plot

It is also possible to adjust the dissipation data of Fig. 4 using a back-extrapolation technique on a square-root time plot, similar to the Taylor method used for interpreting t_{50} values from laboratory 1D incremental consolidation testing. In the root-time plot, the dissipation after the peak caused by redistribution of pore pressure, initially depicts a straight line which can be back-

extrapolated to $t=0$ in order to obtain a modified u_c for the corrected dissipation curve. This value is then used to produce the normalized dissipation curve. The basis of the above correction technique is illustrated in Fig. 7a for measurement locations behind the tip. The principle is the same for measurement locations on the tip except that instead of an increasing pore pressure, the initial pore pressure suddenly drops as discussed previously (Fig. 7b).

The data from Fig. 4 have been plotted in the root time base and are presented in Fig. 8. Direct estimates of t_{50} (50% reduction of excess pore pressure) can be read directly from Fig. 8.

The additional advantage of the root-time method is that the initial straight line portion can be extrapolated to 50% pore pressure reduction if short dissipation periods are used in the field and measured data to longer periods are not available (Fig. 9), or the initial linear slope in the normalized pore pressure root time plot can be analyzed to provide estimates of c_h using the theoretical approach suggested by Teh (1987).

The two correction methods described above will give rise to slightly different normalized dissipation curves since the initial corrected u_c values are, by definition, not the same. The resulting corrected dissipation data for the Strong Pit site are compared in Fig. 10 (the root time plot has been reproduced in log time space for comparison purposes). While the dissipation curves for U less than 25% for both corrections may be different (as would be expected from the different u_i values) at $U=50\%$ the error between the predicted values is relatively small (5% to 10%). In essence, the two correction techniques give similar values for t_{50} . The curves in Fig. 10 do however indicate the importance of the initial pore pressure value at $t=0$ on the normalized form of the dissipation curve.

EVALUATION OF PROPOSED PROCEDURES IN OC SOILS

Basis of comparison

The two correction procedures have been applied to CPTU dissipation data from three University of British Columbia research sites where overconsolidated soils are present in the profile. These extrapolation techniques have also been verified using data presented by some of the major piezocone research centers worldwide (Robertson et al. 1991). The results of comparisons between the log t correction and laboratory derived consolidation parameters are considered below. The root time method is not presented here since the difference between the resulting times from the two methods is only significant for short dissipation periods. All the results presented here are for 50% dissipation of the excess pore pressure and, as suggested by Fig. 10, the difference between the two values of t_{50} is insignificant when taken in context of the overall magnitudes involved.

The horizontal coefficient of consolidation can be evaluated from the corrected CPTU data using any of the available theories. For this study, the method proposed by Teh (1987) has been used. The advantage of this method is that it considers the effect of the rigidity index on the pore pressure dissipation. The c_h value is determined from:

$$c_h = (T^* R^2 I_R^{0.5}) / t_{50} \quad (4)$$

where:

T^* is the Teh and Houlsby (1988) modified time factor

R is the cone radius at the measurement location

I_R is the rigidity index of the soil

Results presented by Danziger (1990) suggest that the Houlsby and Teh (1988) approach provides more consistent c_h estimates than the other available methods.

For comparison with the CPTU interpretation, consolidation coefficients from incremental laboratory oedometer tests are presented. While these laboratory determined c_v values may not be wholly representative of in situ conditions, the results do provide a basis on which the relative CPTU magnitudes can be compared. Field data are available from several international research sites where CPTU derived $c_{h,v}$ values can be compared to the results of large-scale field tests (Robertson et al. 1991).

Laboratory c_v values are determined from each loading stage so that both OC and NC data are available. Furthermore 1D oedometer tests have also been performed on directionally-cut samples so as to obtain estimates of c_h . Hence the c_h/c_v ratio can be estimated from the laboratory tests in order to correct the CPTU c_h values to c_v for direct comparison with the laboratory data.

Geotechnical review of UBC sites considered

The general geotechnical characteristics of the UBC test sites under consideration are presented in Table 1. The soils at Lr. 232 St. are soft sensitive clay silts. At the other two sites the clay silts are non-sensitive with undrained strengths up to 200 kPa.

Table 1 Geotechnical characteristics for UBC test sites considered.

Site	Depth range (m)	PI (%)	OCR	Range of σ_{vo}' (kPa)	Range of σ_{vm}' (kPa)
Lr. 232 St.	1-5	21-30	3-10	16-40	90-205
Strong Pit	1-9	11-20	2-15	16-180	350-500
200 th St.	1-5	20	2-17	16-51	115-300

σ_{vm}' is the maximum past vertical effective pressure from incremental oedometer tests

Results of comparison for UBC sites

Only data from pore pressure locations u_1 and u_2 have been used for the purposes of comparison. The times required for 50% dissipation of the excess pore pressure measured at the u_3 location are prohibitively long and the use of piezocone dissipation tests at this location is not considered by the authors to be of practical interest unless graphical extrapolation techniques can be employed. This is evident from the data in Fig. 6 which highlight the problems associated with the execution and interpretation of dissipation data at locations behind the tip, since the time periods required are very long in low permeability soils. The time for 50% dissipation at the u_1 location is 1100 seconds (18 min.) which increases to 2300 seconds for u_2 and approximately 10 000 seconds for u_3 (the u_3 record has been extrapolated for comparative purposes to obtain the 50% dissipation times).

The coefficients of consolidation from CPTU (c_h) and laboratory oedometer (c_v) tests have been determined as described above. The obtained values are compared in Table 2.

The c_h values from the two pore pressure measurement locations (u_1 and u_2) are in very good agreement with each other. Furthermore, the CPTU values would suggest that the t_{50} dissipation provides an estimate of the c_h corresponding to the OC condition, i.e. the in situ condition of the soil. Considering the results presented by Baligh and Levadoux (1980) who show that, in NC soils, the t_{50} time corresponds to the $(c_h)_{NC}$, it would appear that at degrees of dissipation of 50%, the theory provides estimates of the coefficient of consolidation relevant to the in situ stress history condition, $(c_h)_{OC}$. In addition to this, it is apparent from the types of dissipation curve considered, that stress history may be an important factor to be considered when interpreting CPTU dissipation data. As suggested by Lutenegeger and Kabir (1987), the dissipation curve may be useful as a stress history indicator.

Table 2 Comparison of field and laboratory coefficients of consolidation

Site	Filter location	$(c_h)_{OC}$ from CPTU (cm^2/s)	c_v from oedom. (cm^2/s)		c_h/c_v (oedom)
			OC	NC	
Lr.232	u_1	0.002-0.005	0.006-0.1	0.0005-0.001	2-3
	u_2	0.005-0.016			
Strong	u_1	0.007-0.004	0.002-0.005	0.0006-0.001	1-2
	u_2	0.01-0.006			
200 St.	u_1	0.014-0.047	0.05-0.18	0.001-0.03	1.5-2
	u_2	0.045-0.054			

Evaluation of data from international sites

Dissipation tests using the cone penetrometer have been carried out in a variety of overconsolidated soils worldwide. The data have been summarized by Robertson et al. (1991). Typical results highlighting the different types of pore pressure response described above are presented below.

The results of a u_1 dissipation in moderately overconsolidated Haga clay (NGI 1990) are presented in Fig. 11a. The pore pressure of about 875 kPa generated during penetration decreases immediately to 480 kPa after stopping penetration. Thereafter, the dissipation occurs as the excess pore pressure close to the cone dissipates laterally into the surrounding soil. This is similar to the Type II response described earlier. The lightly overconsolidated glacial till at Cowden (OCR = 3) provides an example of a Type III dissipation at the u_2 location (BRE/NGI 1985). The positive penetration pore pressure of 835 kPa increases after stopping penetration until a maximum value of around 980 kPa is reached some 10 seconds later (Fig. 11b). Correcting the dissipation curves for the maximum values and using the new zero times, as described earlier, allows the modified dissipation curve to be interpreted to obtain the OC coefficient of consolidation. The data from Brent Cross (OCR > 30) are also characteristic of a Type III response for the u_2 location (Fig. 11c).

The Richards Island silty clay (OCR = 8) dissipation results presented by Campanella et al. (1986) are typical of a Type III response for the u_3 location (Fig. 11d). For this case the penetration pore pressure is only marginally higher than the in situ equilibrium value, u_o - if it had been lower than u_o , then this would have been a Type IV dissipation response, similar to the result obtained at 200th St. in the Lower Mainland of British Columbia (Fig. 11e) where the penetration

pore pressure at the start of the dissipation ($u_i = -9 \text{ m H}_2\text{O}$) is less than the equilibrium value ($u_o = 1.6 \text{ m H}_2\text{O}$) in this moderately overconsolidated clay ($\text{OCR} = 10$). As the permeability of the soil increases, the Type IV response changes to a Type V response, where the negative excess generated during penetration dissipates directly to the equilibrium value. This type of response has been recorded in dilative silts and sandy silty clays (Fig. 11f).

All the OC dissipation data discussed above were presented by Robertson et al. (1991), having been interpreted using the above correction techniques. This data and other results for OC soils presented by Robertson et al. but not discussed here, are plotted in Fig. 12, using the c_h - t_{50} nomograph. In Fig. 12 the CPTU dissipation data are compared to c_h values obtained from laboratory consolidation tests. Data have been reported in the literature for a limited number of sites, where c_h has also been determined from back analysis of field performance. The published values of c_h obtained from field performance are generally larger than laboratory values (Robertson et al. 1991, Jones and Rust 1993) and provide better agreement with the theoretical range obtained from the Teh and Houlsby (1991) solution.

CONCLUSIONS

The paper has presented a classification for the types of dissipation response recorded in overconsolidated fine grained soils during piezocone penetration tests. The different dissipation responses have been explained in terms of the unloading and redistribution that occurs as the soil undergoes lateral displacement to permit advancement of the cone tip and shaft. The singularity at the shoulder of the cone tip gives rise to the unloading as the soil no longer

undergoes cavity expansion type deformation, but rather shearing along the shaft-soil interface. The soil stiffness and strength, which can be considered to be dependent on the consolidation state of the soil, control the degree of unloading and redistribution, and consequently the extent of the modified pore pressure response (compared to the Type I response associated with normally consolidated soils). Also important in controlling the type of response is the location of the porous element used for measuring the pore pressure variation with time.

Two data manipulation methods have been presented for correcting the different types of pore pressure response in OC soils. In one method, the data in log t space is normalized to the maximum post-penetration value and the time at which this maximum value occurs. In the second approach, a root-time correction and extrapolation is used to correct the field data (see Appendix 1). Once either of the above modifications has been made, the data can be interpreted using standard techniques available for normally consolidated soils. The Teh and Houlsby (1988) method has been used in the text to interpret the data using the nomograph presented by Robertson et al. (1991).

The c_h values from CPTU generally fall into the range suggested by the laboratory data for the overconsolidated state for the in situ tests at the wide range of sites considered. The correction method proposed here appears to provide consistent estimates of c_h based on a theoretically acceptable technique.

It would also seem reasonable to recommend a preferential pore pressure measurement location in terms of performing CPTU dissipation tests. The u_3 location involves prohibitively long dissipation times to be of practical use. On the other hand, the u_1 position would appear to be a good location due to the higher rate of dissipation involved (since this is also function of the square of the cone radius). However, since the u_2 position is required for correction of pore

pressure end area effects, this would appear to be a reasonable compromise. This would concur with the recommendation of Robertson et al. (1991) who suggest that more consistent and repeatable c_h determinations are obtained at the u_2 location than either u_1 or u_3 (with respect to both field and laboratory data).

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APPENDIX 1 - Correction Procedure for Dissipation Response Types II to V

A1. Evaluate dissipation record in terms of pore pressure response with time and, based on pore pressure measurement location (u_1 , u_2 or u_3), determine type of response according to classification scheme presented above.

A2. If the filter is on the cone tip (u_1) and the pore pressure drops after stopping penetration (Type II response), determine point A (Fig. 5a) at which pore pressure starts to dissipate after unloading has occurred.

If the filter is located behind the cone tip (u_2 or u_3), determine point at which post-penetration increase stops and pore pressure dissipation commences (point B, Fig. 5b and 5c). For a Type III or IV response, the dissipation can be interpreted using the post peak decay (Fig. 5b and 5c), whereas for a Type V record (Fig. 5c), the post-penetration rise to u_0 can be treated as an inverse dissipation to provide estimates of the coefficient of consolidation.

A3. Using the data point (u_c and t_c) determined above in A2, redraw the corrected pore pressure dissipation curve (where u_c is now the maximum pore pressure and t_c is the zero time reading) to obtain a standard dissipation plot typical for normally consolidated soils. The data can be plotted either in log time or root time and the values of t_{90} or t_{50} read off. The values thus obtained can be used to calculate the coefficient of consolidation using the appropriate theory.

APPENDIX 2 - Classification used for Overconsolidation State

The following classification is used here for description of the soil overconsolidation state:

Stress History	<u>Description</u>
OCR=1	Normally consolidated (NC)
1<OCR=4	Lightly overconsolidated (LOC)
4<OCR=10	Moderately overconsolidated (MOC)
10<OCR=25	Heavily overconsolidated (HOC)
OCR>25	Very heavily OC (VHOC)

$$\text{OCR} = \text{Overconsolidation Ratio} = \sigma'_{vm} / \sigma'_{vo}$$

where:

σ'_{vm} is the maximum vertical effective pressure the soil has experienced

σ'_{vo} is the vertical effective stress presently acting in the ground on the soil

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Fig. 1 Terminology used for indicating location of pore pressure measurement

Fig. 2 Typical dissipation records for NC fine grained soils

Fig. 3 Pore pressure dissipation in stiff moderately OC silty clay (Strong Pit)

Fig. 4 Normalized pore pressure dissipation for Strong Pit clay

Fig. 5 Idealized pore pressure dissipation response in OC fine grained soil for different measurement locations

Fig. 6 Normalized dissipation log-time plot for data in Fig. 3 corrected for unloading (u_1) and redistribution (u_2 and u_3)

Fig. 7 Details of the root-time method for evaluating pore pressure dissipation data in OC soils according to response type: a) Type II response, b) Type III to IV responses

Fig. 8 Measured dissipation data for stiff clay in root-time plot

Fig. 9 Root-time extrapolation for short duration dissipation records

Fig. 10 Comparison of normalized dissipation curves after applying proposed log-time and root-time corrections

The first part of the report discusses the general situation of the country and the progress of the work done during the year.

The second part of the report deals with the various projects and schemes which have been carried out during the year.

The third part of the report contains a list of the names of the persons who have been employed during the year.

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Fig. 11 Published dissipation data for OC fine grained soils illustrating proposed classification scheme: a) Type II response in Haga clay; b) Type III response in glacial till at Cowden; c) Type III response in London clay at Brent Cross; d) Type III response in OC silty clay at Richards Island; e) Type IV response at 200th St., and f) Type V response in stiff sandy silty clay

Fig. 12 Comparison of t_{s0} from corrected dissipation data in OC soils with c_h from laboratory consolidation tests using dissipation interpretation nomograph presented by Robertson et al. (1991)

The first part of the document discusses the importance of maintaining accurate records of all transactions. It emphasizes that every entry should be supported by a valid receipt or invoice. This ensures transparency and allows for easy verification of the data.

In addition, it is crucial to review the records regularly to identify any discrepancies or errors. This proactive approach helps in resolving issues before they become significant problems.

The following table provides a summary of the key points discussed above.

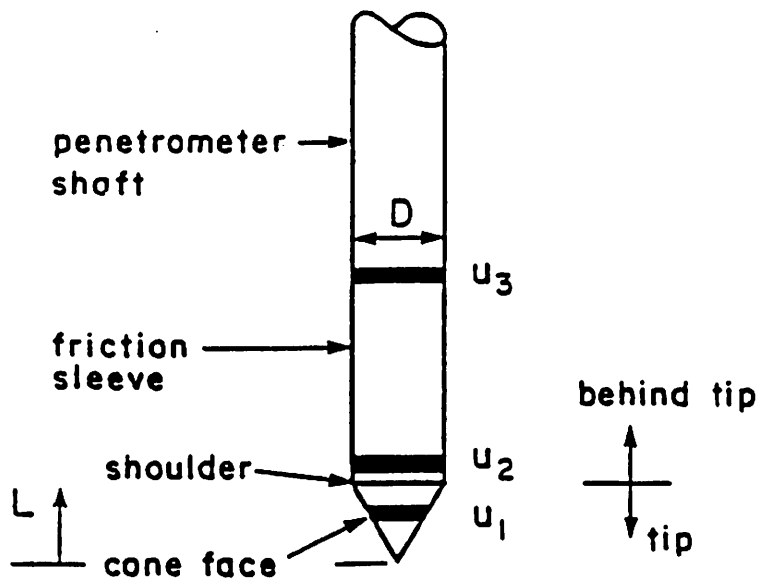


Fig. 1 Pore pressure measurement locations considered

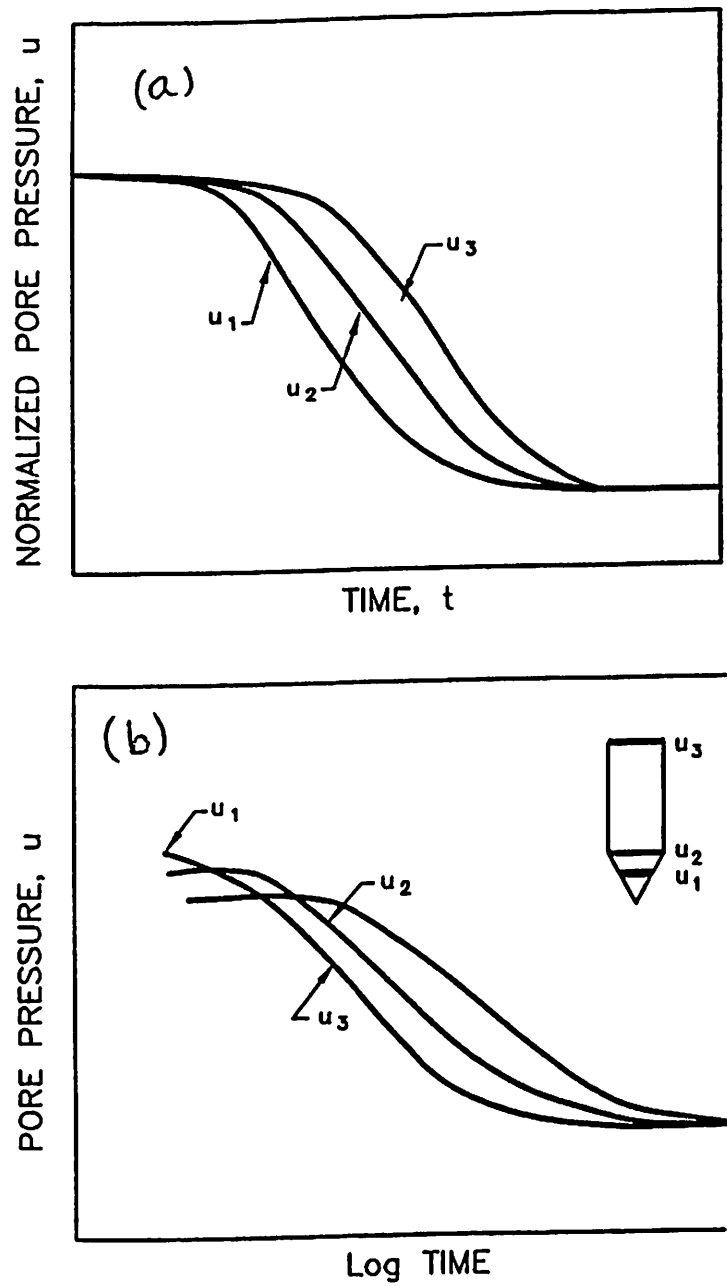


Fig. 2 Typical dissipation records for NC fine grained soils

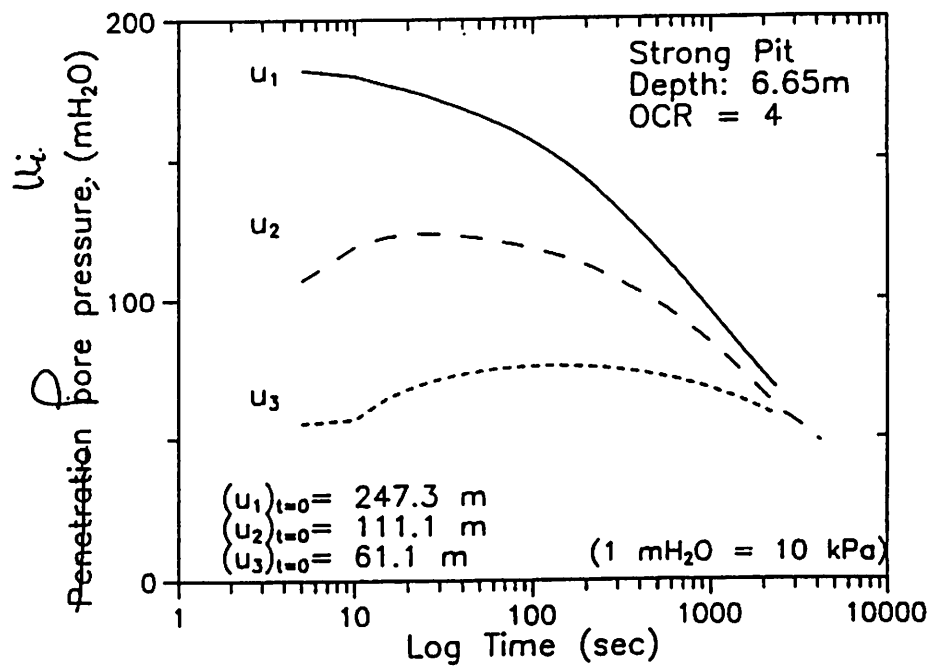


Fig. 3 Pore pressure dissipation in lightly OC clay

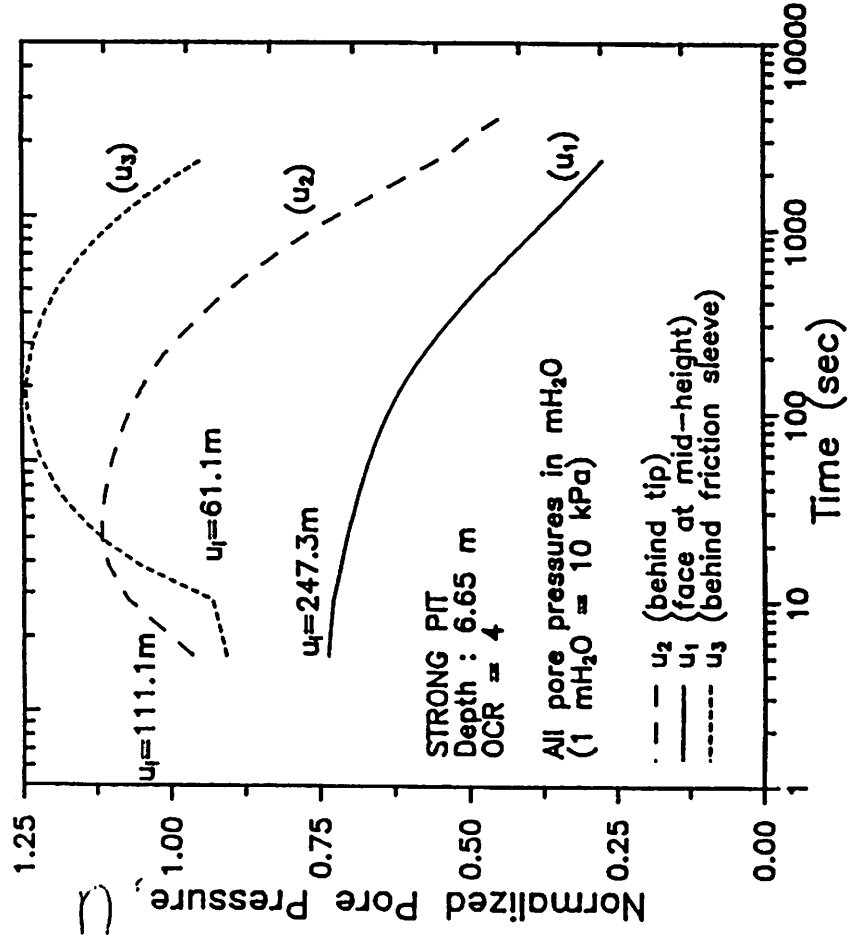


Fig. 4 Normalized pore pressure dissipation for Strong Pit clay

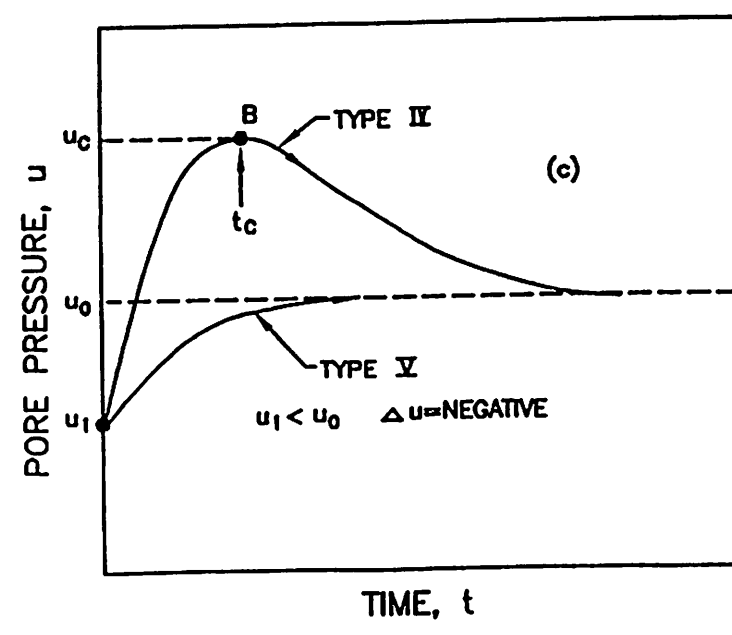
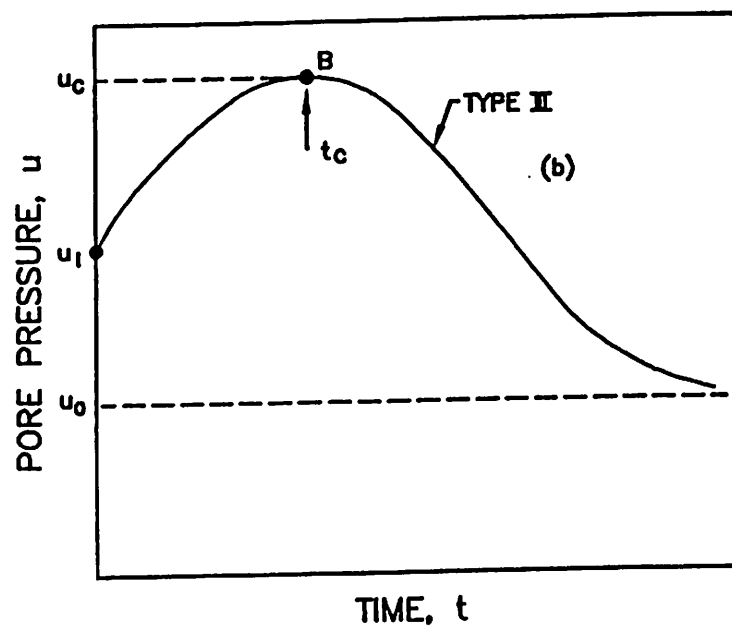
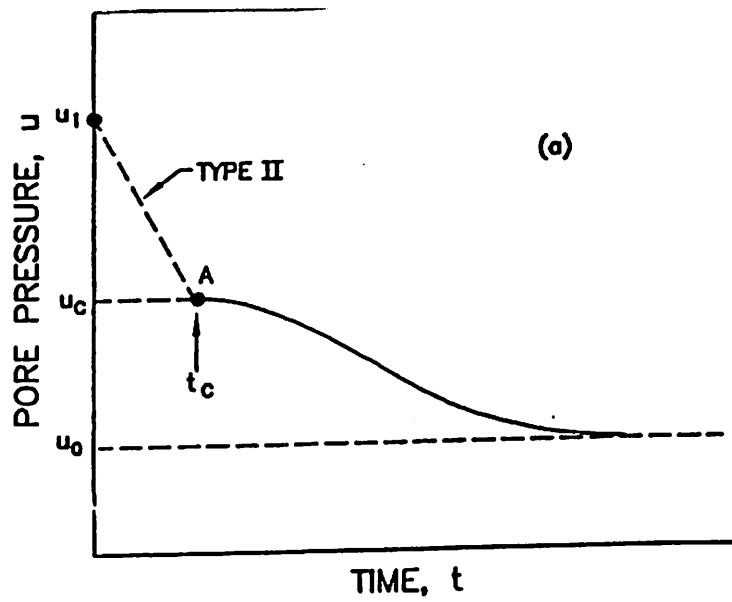


Fig. 5 Idealized pore pressure dissipation response in OC fine grained soil for different measurement locations

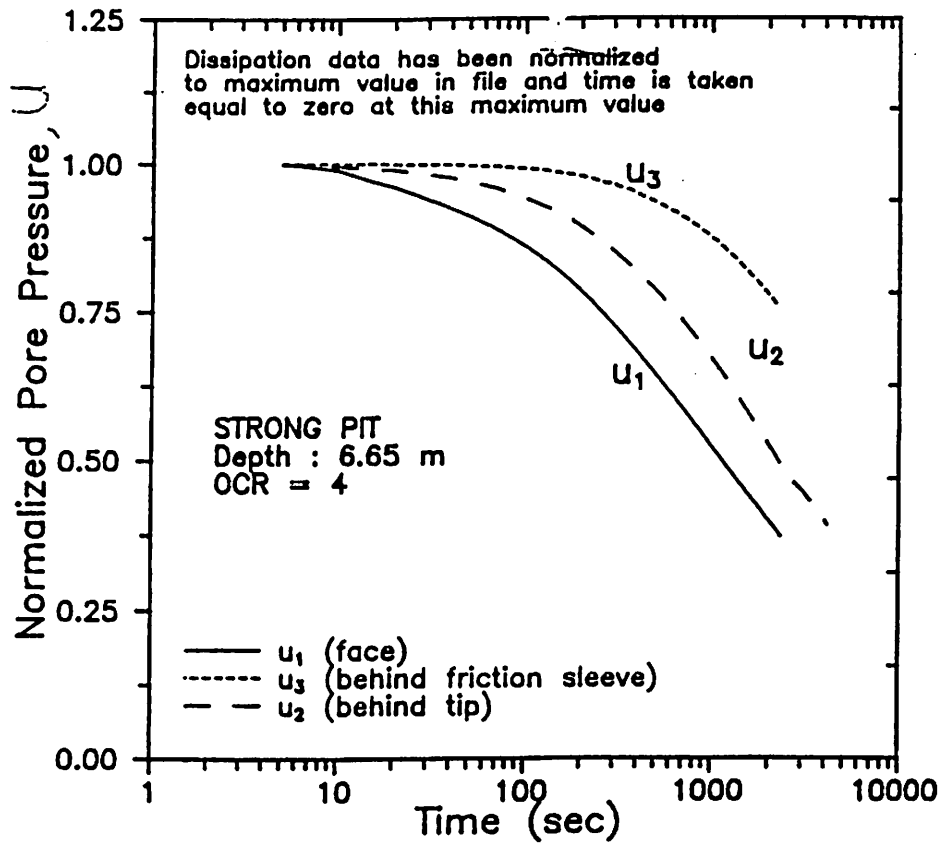


Fig. 6 Normalized dissipation log-time plot for data in Fig. 3 corrected for unloading (u_1) and redistribution (u_2 and u_3)

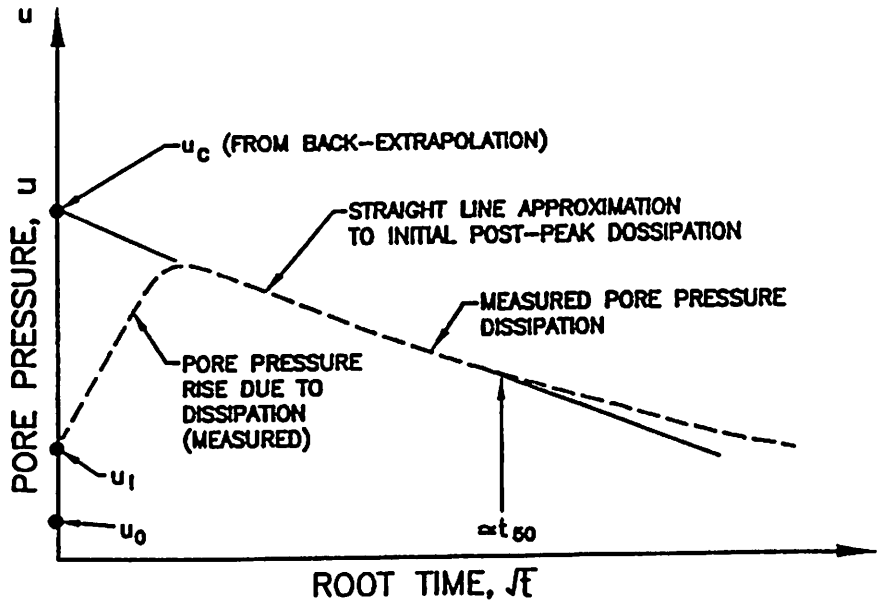
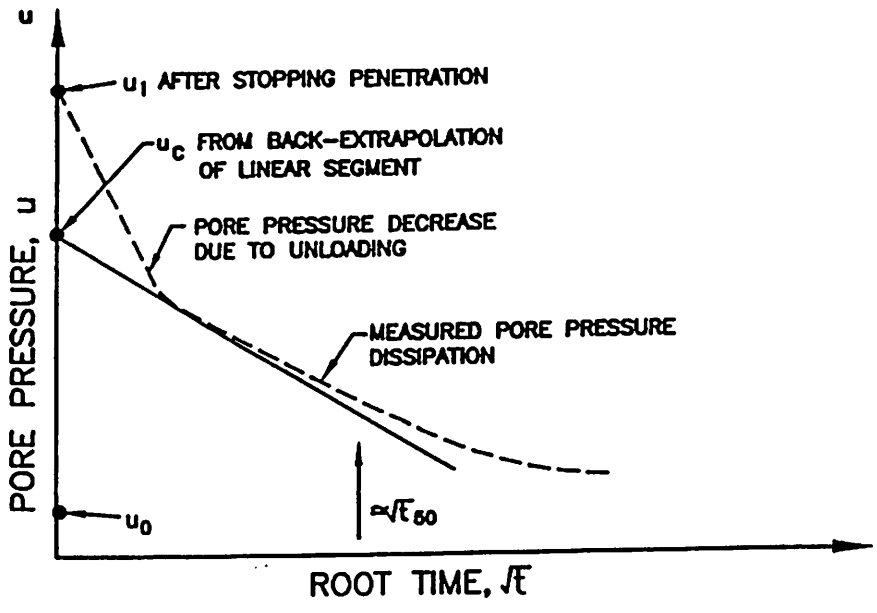


Fig. 7 Details of the root-time method for evaluating pore pressure dissipation data in OC soils according to response type: a) Type II response, b) Type III to IV responses

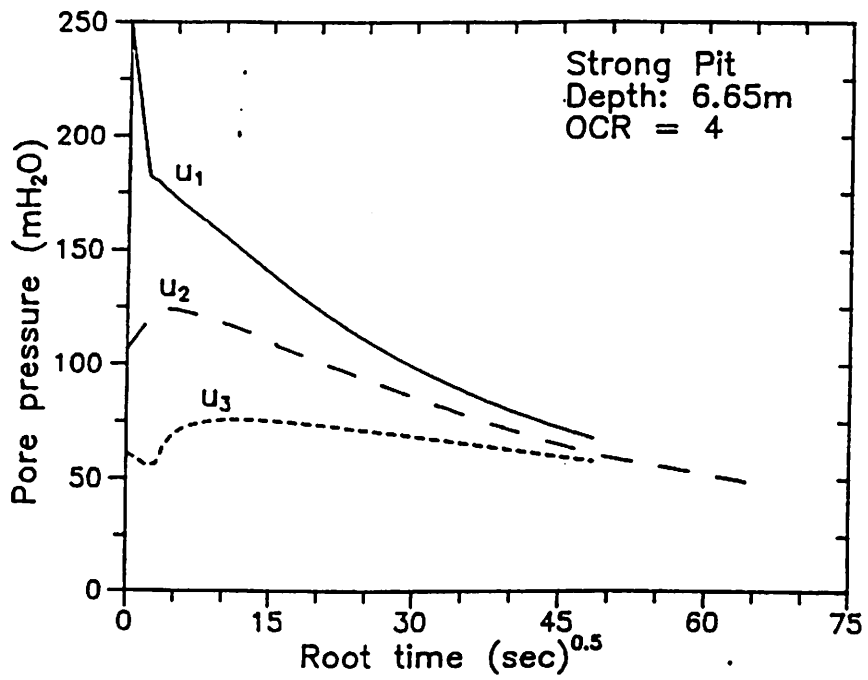


Fig. 8 Dissipation data from stiff clay in root time plot

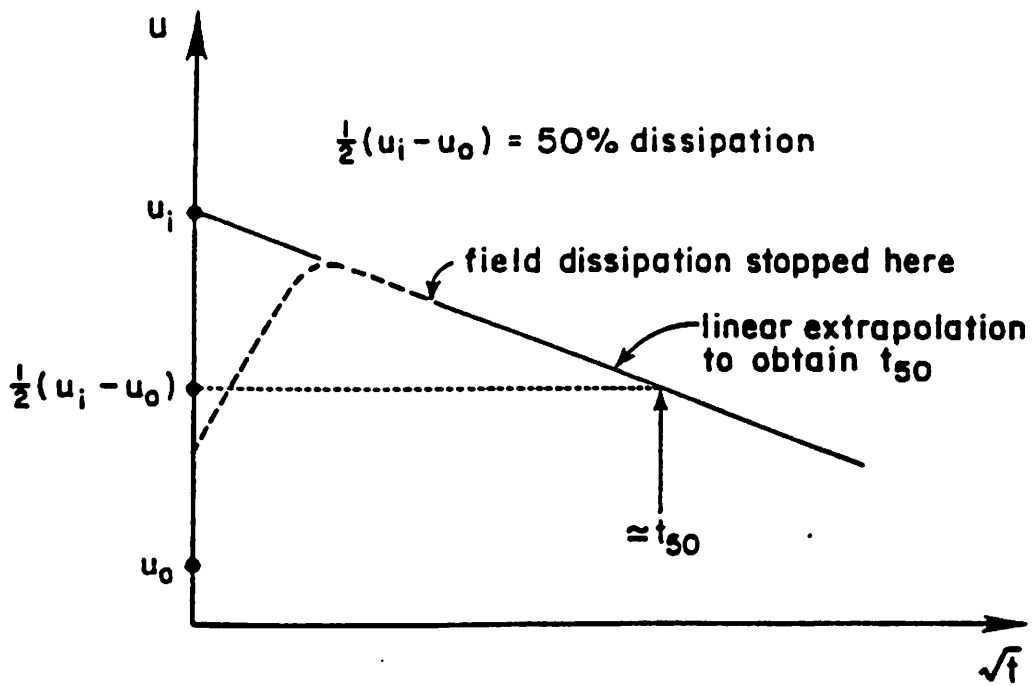


Fig. 9 Root-time extrapolation for short duration dissipation records

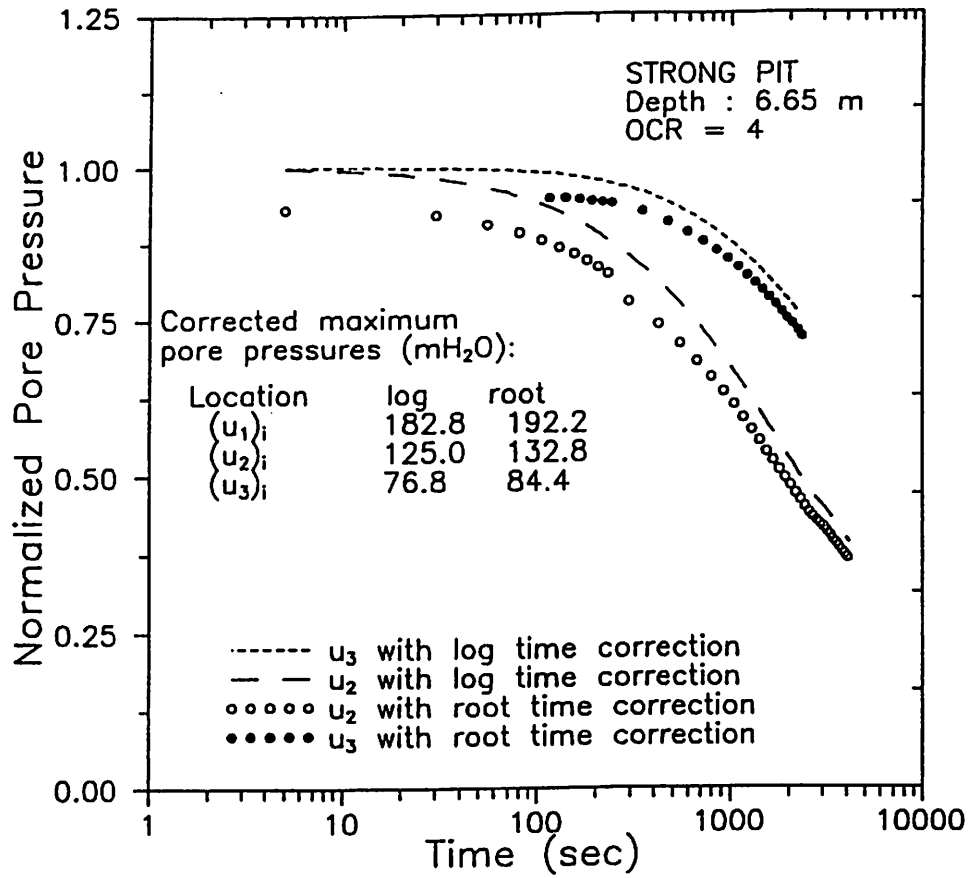
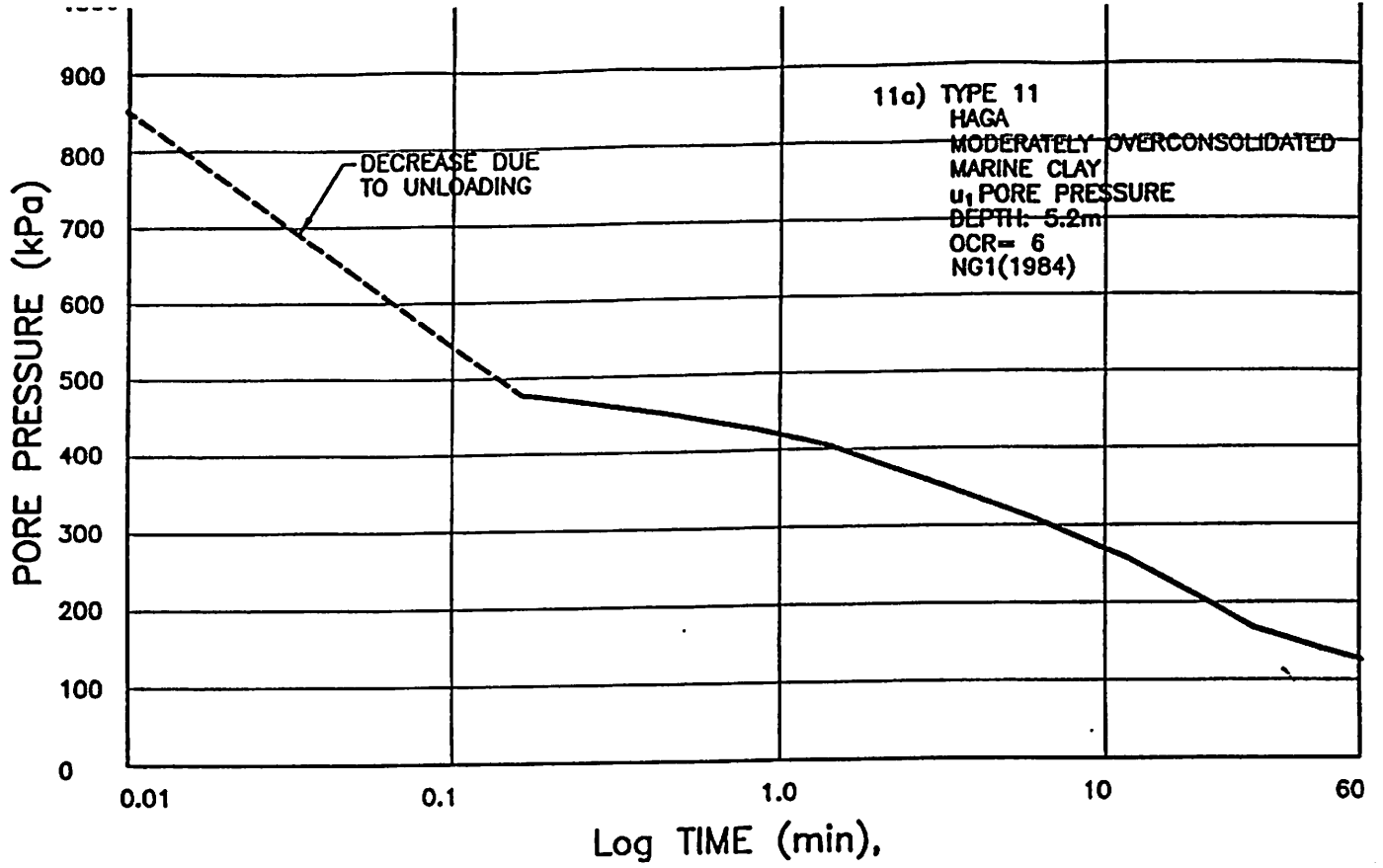


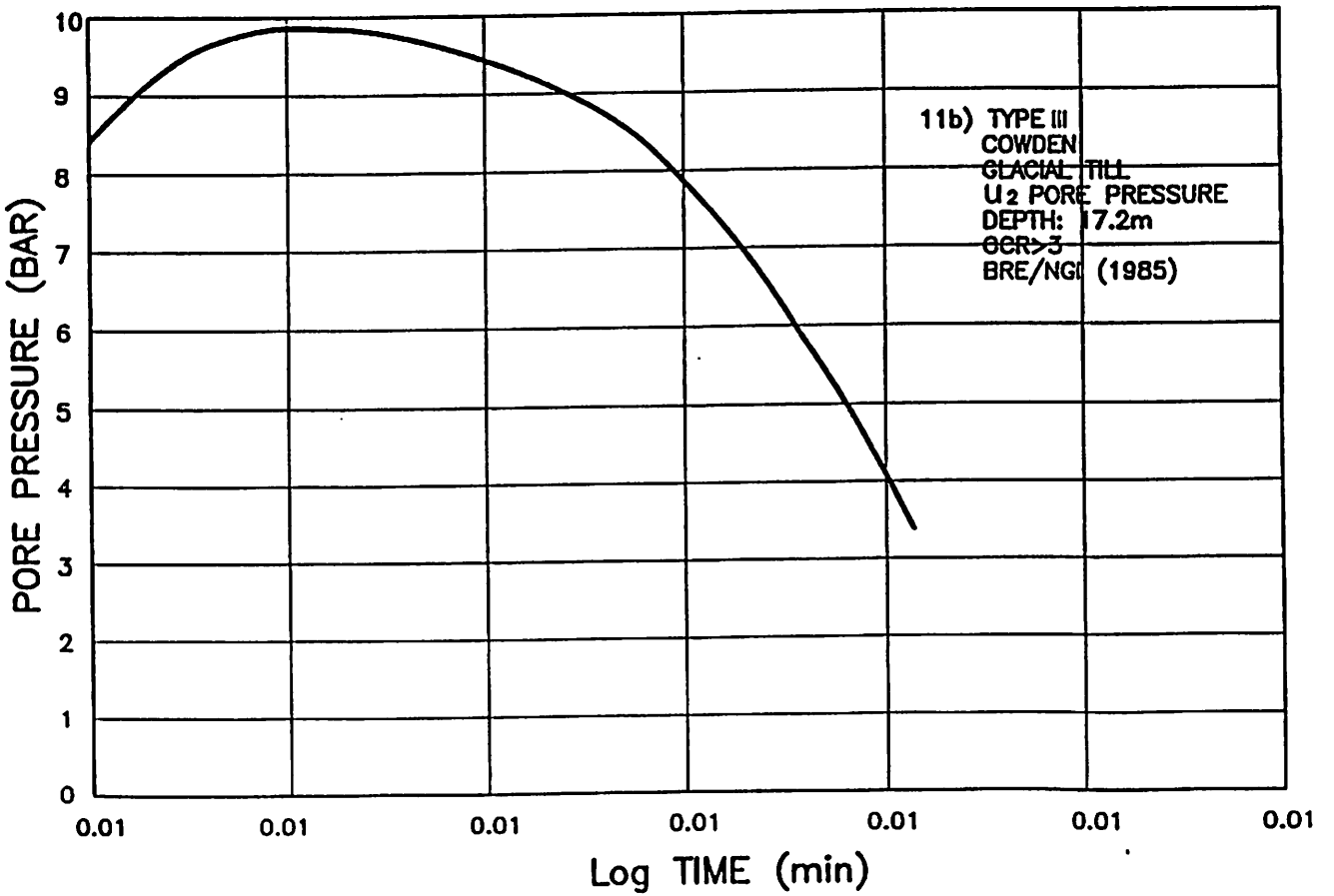
Fig. 10 Comparison of normalized dissipation curves after applying proposed log-time and root-time corrections

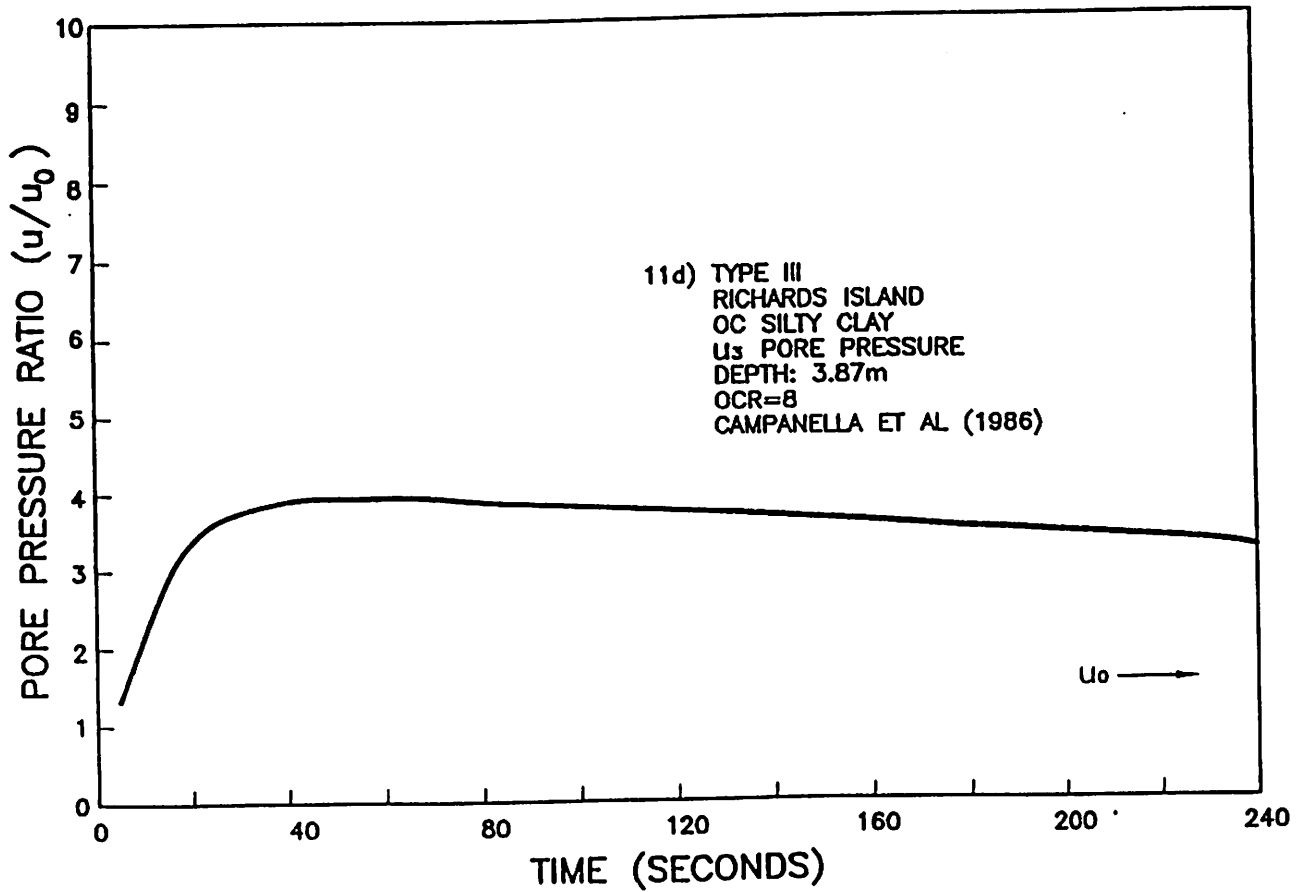
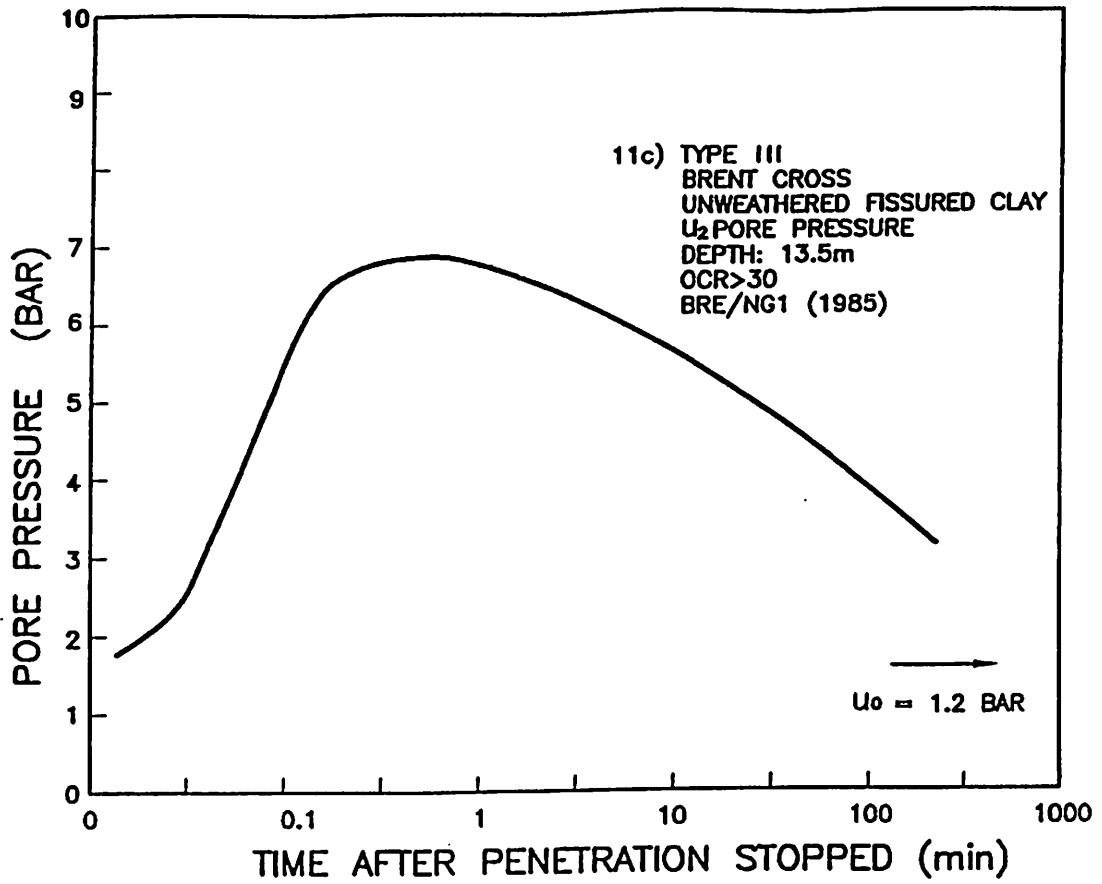
Fig. 11 Published dissipation data for OC fine grained soils illustrating proposed classification scheme: a) Type II response in Haga clay; b) Type III response in glacial till at Cowden; c) Type III response in London clay at Brent Cross; d) Type III response in OC silty clay at Richards Island; e) Type IV response at 200th St., and f) Type V response in stiff sandy silty clay

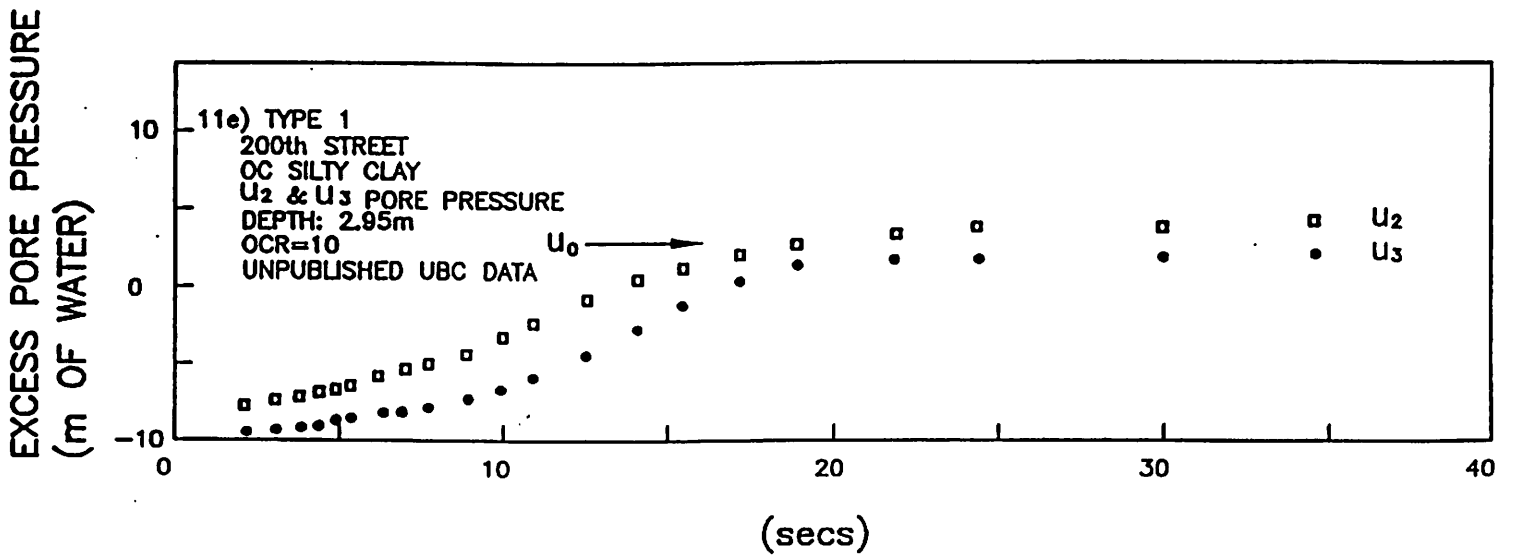
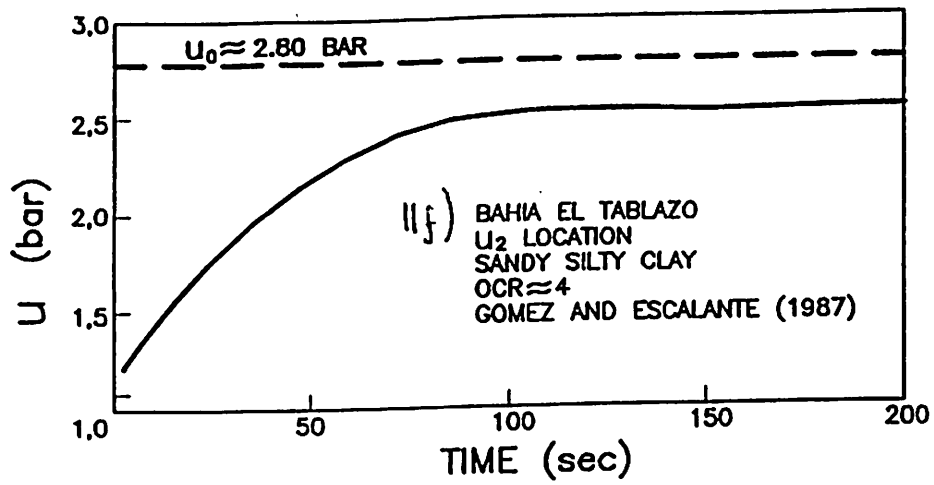
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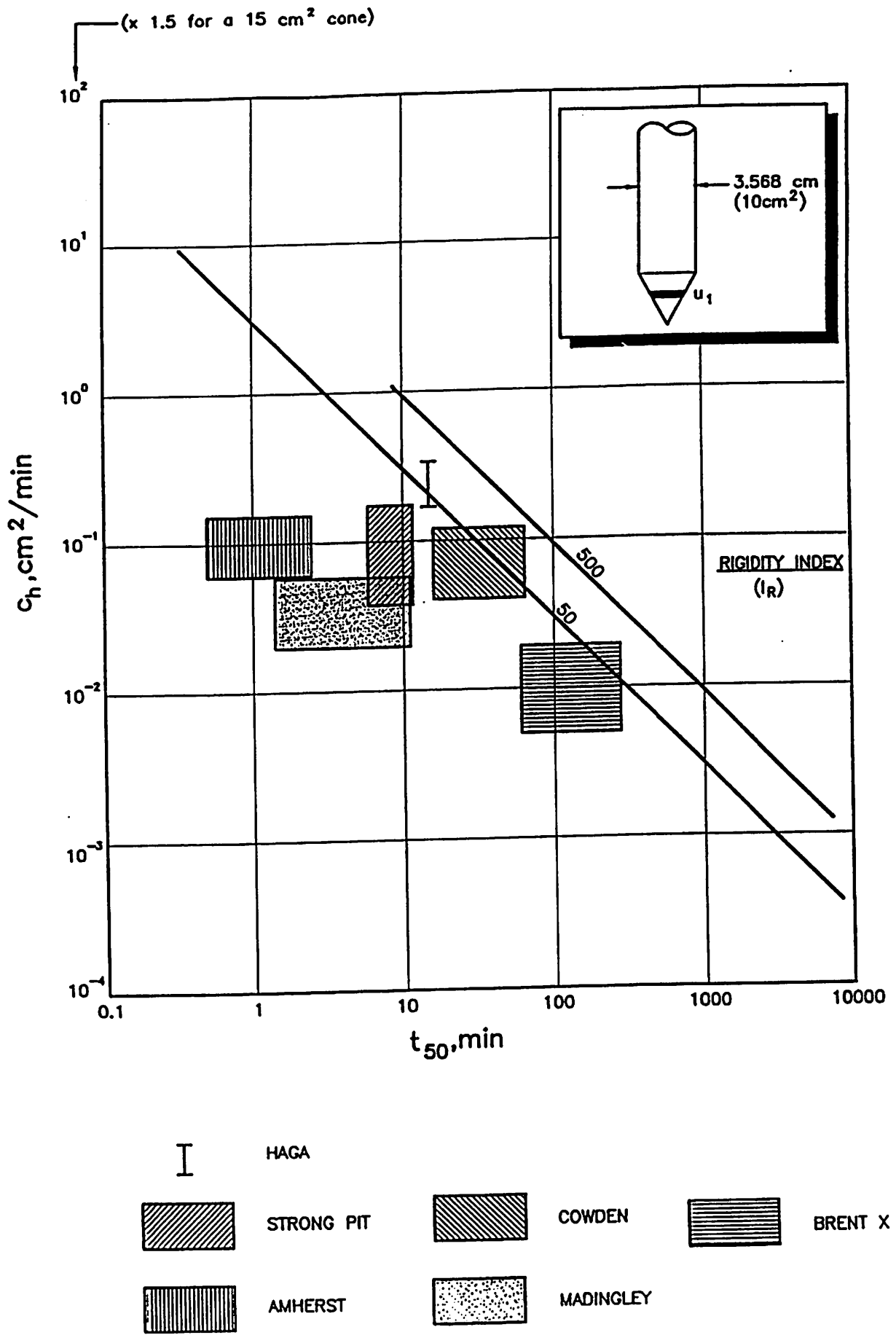


Fig. 12 Comparison of t_{50} from corrected dissipation data in OC soils with c_h from laboratory consolidation tests using dissipation interpretation nomograph presented by Robertson et al. (1991)