NEOGENE PALEOGEOGRAPHY OF THE WESTERN UNITED STATES

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ABSTRACT

This paper presents a brief introduction to the Miocene and Pliocene paleogeography of the western portion of the United States including California, Oregon, Washington, Nevada, and the western portions of Arizona and Idaho. The maps are composites and attempt to illustrate maximum marine inundation during Miocene and Pliocene time.

Paleogeography as used in the text is defined as the study of the distribution of land, sea, air, floral and faunal provinces on the Larth's surface at a specified point in time.

During Miocene time, until about 19 my ago, the Basin and Range province was dominated by intermediate type volcanism. Approximately 16 my ago, basaltic volcanism commenced contemporaneously with normal faulting and produced the Basin and Range topography associated with the region today.

West of the Basin and Range, the Sierra Nevada was an andesitic volcanic arc in which arc volcanism was progressively truncated from south to north. This cessation was coincident with the northward migration of a triple junction which resulted from the collision of the East Pacific Rise with North America. This Cordilleran volcanic arc continued north into Oregon and Washington where it is still active. In California west of the Sierra Nevada, a marine basin was disrupted by the previously mentioned collision and subsequent offset along the San Andreas fault.

Western Oregon and Washington were dominated in Miocene and Pliocene time by an andesitic volcanic arc. West of the arc, the Pacific Ocean extended inland only slightly farther east than today's shoreline.

East of the arc, the middle and late Miocene outpouring of the Columbia River Group basalts was the dominante event. This was followed by extrusion of the Snake River Plain volcanic rocks which show a progressive decrease in age of inception eastward to Yellowstone where they terminate.

INTRODUCTION

This paper presents a thumbnail sketch of the Miocene and Pliocene paleogeography of the United States west of 1110 latitude. Neogene deposition and paleogeography are briefly discussed and illustrated on the accompanying maps. The maps are schematic interpretations of an average "paleogeography" for each time slice and attempt to illustrate the maximum marine in undation. The Miocene map encompasses events between 22 and 5 million years; the Piliocene map encompasses events between 5 and 2 million years.

The Miocene map generally illustrates late Miocene paleogeography. In California rocks of the late Miocene represent the maximum marine Neogene inundation, while in Oregon and Washington maximum inundation occurred during the early Miocene. The maximum inundation in Oregon and Washington is therefore not represented on the Miocene map. The magnitude of this problem is not as great as it first appears, for the late Miocene inundation in Oregon and Washington, which is represented on the map, reoccupied to a lesser degree the same embayments as the early Miocene inundation. This lack of regionally synchronous events can be overcome by generating maps for narrower time Preparation of this report, however slices. permitted synthesis of only two maps for a time span perhaps best illustrated by at least four maps.

The Pliocene map, encompassing a shorter period than the Miocene map, compresses a less complex geologic history into a single plane. Pliocene paleogeographic features have, with some modification, persisted to the present.

The Miocene and Pliocene maps have been generated using King and Beikman (1974) as a geologic base. With properly scaled reproductions, the paleogeographic maps of this report may be superimposed upon King and Beikman's map allowing direct identification of the geologic terrains forming the base for the paleogeographic interpretations.

This approach produces a simplistic view of paleogeography: a view that ignores palinspastic reconstructions. Detailed studies of the San Andreas fault system have delineated major offset features indicating approximately 190 mi (315 km) of right slip in the last 15 million years (Nilsen and Link, 1975).

On the Miocene map, submarine fans truncated by the fault are shown entering the area of the modern San Joaquin Valley from the west. Across the fault to the west, alluvial fans truncated by the fault are shown with an eastern source. Restoration of the post-Miocene offset will restore both the submarine fans and alluvial fans to their respective source terrains (Huffman, 1972).

The paleogeographic consequence of other dislocated regional geologic terrains is less well understood. Paleogeomagnetic data suggests Cenozoic rotation of volcanic terrains in both western Oregon (Simpson and Cox, 1978) and southern California (Kamerling and Luyendyk, 1977; Kamerling and others, 1978). Simpson and Cox (1978) interpret paleomagnetic data to show that Eocene to Miocene volcanic rocks in western Oregon have been progressively rotated clockwise as much as 65°. Based on that data, Simpson and Cox proposed several tectonic reconstructions which encompassed all of the western Oregon and Washington Coast Ranges. Unpublished paleogeomagnetic data (M. E. Beck, 1979, personal comm.; Ray Wells, 1979, personal comm.) show that Eocene volcanic terrains in southwestern Washington have rotational histories different than age equivalent rocks in western Oregon. Such data requires that each volcanic terrain with a unique rotational history be treated as a separate tectonic block. As yet we do not know how many of these separate tectonic blocks we are dealing with, nor do we know either the nature of or location of the structural boundaries between blocks with unique tectonics rotational histories. In light of these unresolved palinspastic problems we decided a non-palinspastic reconstruction of the Meogene paleogeography would be more useful especially when used in conjunction with the geologic base of King and Beikman (1974).

PALEOGEOGRAPHY

The Cenozoic geology of western North America formed around a framework of Paleozoic and Mesozoic plutonic and metamorphic terrains (see volume I and II of this series). Today, these terrains form uplifted areas such as the Sierra Nevada, Klamath Mountains, Blue Mountains, Colorado Plateau, Rocky Mountains of Idaho, Washington and southern Canada, the North Cascade Mountains, and the Coast Mountains of British Columbia. Neogene deposits derived from these pre-Cenozoic terrains together with intra-basinal derived deposits (mostly volcanic rocks) accumulated in depositional basins such as 1) the Basin and Range; 2) the Columbia Plateau and Snake River Plain; 3) the Sierra Nevada; 4)

the Cascade Mountains; 5) the Puget-Willamette Lowland; 6) the Great Valley of California; and 7) the Coast Ranges and continental shelves of California, Oregon and Washington. These geologic provinces are outlined in figure 1. The Miocene and Pliocene depositional patterns and paleogeography are illustrated on figures 2 and 3.

BASIN AND RANGE

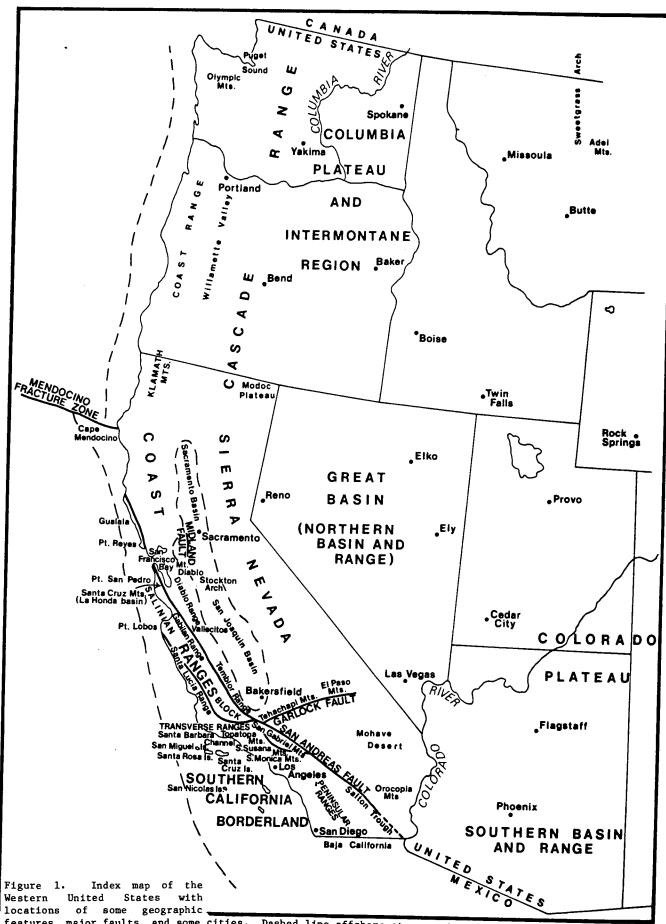
The geologic province east of the Sierra Nevada and west of the Rocky Mountains and Colorado Plateau is the Basin and Range province. This province is bounded on the north by the Columbia Plateau and Snake River Plain, and merges on the south with the Mojave Desert. The Basin and Range province gets it name from the many north-south aligned basins separated by mountain ranges.

The early Miocene rocks of the Basin and Range province are intermediate rhyolitic volcanics and interbedded non-marine deposits emplaced on an old erosion surface of low to moderate relief. This outpouring of large volumes of lave was not accompanied by any marked increase in tectonism with the exception of caldera collapse resulting from volcanic eruptions (McKee, 1971).

The intermediate rhyolitic phase of volcanism ceased about 19 m.y. ago with the area covered by essentially unfaulted flows. About 16 m.y. ago basaltic volcanism began coincident with high-angle normal faulting (McKee, 1971; Stewart and others, 1977).

The normal faulting which produced a horstgraben terrain resulted in 100-112 mi (160-180 km) of extension of the Basin and Range province along a northwest-southeast axis (Proffett, 1977). The grabens became sites of interior drainage in which were deposited clastic wedges of fluvial, alluvial and lacustrine origin. In southern Arizona and southern Nevada a few of these basins accumulated thick sequences of evaporite deposits (Eberly and Stanley, 1972). Based on newly described marine limestones near the Grand Wash Cliffs, Blair (1978) projects the Gulf of California north along the present course of the Colorado River during the late Miocene. By the late Mioceneearly Pliocene the Colorado River had formed (McKee and McKee, 1972). Discharge from the Colorado River into the marine embayment of the Gulf of California gradually filled the estuary and caused the river mouth to migrate southward toward its present position.

The pattern of Basin-Range extension and non-marine sedimentation persisted through the Pliocene and Pleistocene. As the depositional basins filled the drainage systems began to coalesce and exterior drainage developed. Increased rainfall during the late Pleistocene glacial cycles resulted in formation of large lakes.



features, major faults, and some cities. Dashed line offshore at approximate edge of the continental shelf. After Nilsen and McKee (this volume).

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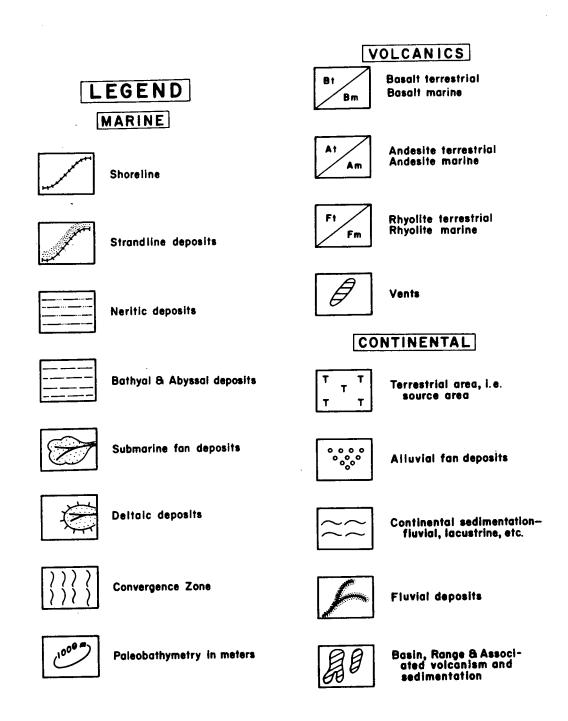
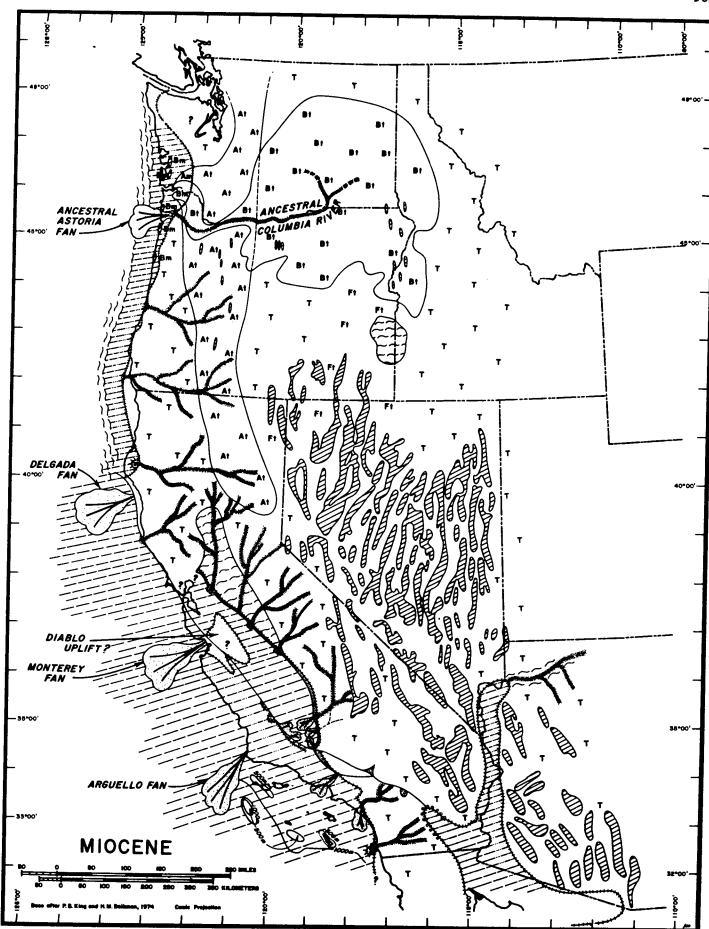


Figure 2. Miocene paleogeographic map of the western United States.



Post-glacial evaporation of these lakes resulted in the complete disappearance of some and the reduction of others to much smaller, more saline lakes.

COLUMBIA PLATEAU AND SNAKE RIVER PLAIN

North of the Basin and Range province east of the Cascade Mountains, nonmarine Neo-gene rocks accumulated in the Columbia Plateau and Snake River Plain province. An early Neogene landscape of rolling volcanic hills drained westward through the ancestral Columbia River system. During middle Miocene the rolling topography was quickly buried by basalt flows of the Columbia River Basalt Group (Waters, 1961) which flowed from fissures and covered much of central Washington and Oregon (Thayer, 1957; Taubeneck, 1970). Coeval lavas of the Steens Basalts covered the Snake River Plain in southeastern Oregon and southern Idaho (Walker, 1970). Younger silicic volcanics of the Snake River Plain group erupted across southern Oregon, northern Nevada and southern Idaho. The Snake River Plain silicic volcanics show a progressive age decrease from west to east; Suppe, Powell, and Berry (1975) suggest this eruption pattern may represent a hot spot track or a propagating crustal crack. present locus of the "hot spot" is Yellowstone National Park.

The Miocene climate of the Columbia Plateau and Snake River Plain was considerably more moist than that of the region today (Axelrod, 1968). Consequently, the processes of weathering, erosion, and deposition rapidly transformed the barren lava covered landscape into an environment with a rich soil, abundant vegetation, lakes and streams. The lakes and streams occupied structural lows formed by late Miocene-Pliocene plateeau subsidence and crustal folding. As the rivers eroded headward the upstream lakes were drained. Subsequent lava flows poured into these river canyons resulted in a complex pattern of plateau lavas, fluvial and lacustrine sediments, and intercanvon lavas (see Swanson, this volume). In cases where the intercanyon lavas only partially filled the canyons the rivers reincised themselves through the intercanyon flows.

During the Pliocene lavas continued to flow across portions of the Columbia Plateau and Snake River Plain with decreasing volume and geographic extent. Uplift of the Cascade Mountains brought increasing amounts of detrital sediment into the western part of the province.

Pleistocene modification of the lava plains of eastern Oregon and Washington and southern Idaho include large expanses of lakes, continued incision and widening of river valleys, and glaciation.

SIERRA NEVADA

The high mountains of eastern California, known as the Sierra Nevada, consisting of Mesozoic plutons and Paleozoic and Mesozoic metamorphic rocks separate the structural lows of the Basin and Range province to the east and the Great Valley to the west. Northward the Sierra Nevada is off-set to the west in the Klamath Mountains (Davis, 1966; Irwin, 1966). The Cenozoic volcanics of the Cascade Mountains overlap the edges of both the northern Sierra Nevada and the eastern Klamath Mountains and extend northward into Oregon and Washington. The southern end of the Sierra Nevada is truncated by the Garlock Fault system. During the Miocene and early Pliocene, the Sierra Nevada was an area of low to moderate relief. Rivers drained westward into a marine basin now occupied by the Great Valley of California (Axelrod, 1956).

Neogene volcanic activity in the Sierra Nevada was progressively terminated in a south to north direction. Middle Miocene volcanic activity extended from south of the Garlock Fault northward throughout the province. By the late Miocene volcanic activity extended northward from a position north of the Garlock Fault. In the early Pliocene, the southern terminus of Sierra Nevada volcanism was on the latitude of San Francisco Bay. The present volcanic activity is limited to areas north of about 40° latitude (Snyder and others, 1976). Atwater (1970) attributed this pattern of volcanism to passage of plate boundaries (see Dickinson, this volume).

The uplift of the Sierra Nevada to the present elevation of over 14,000 feet probably began during Pliocene time (Christensen, 1966). The tilting of the range was progressive with about one-third of the present tilt developed between 9.5 and 18 m.y. ago and the latter two-thirds taking place in the last 9.5 m.y. (Noble and Slemmons, 1975). Based on floral data, the Sierra Nevada was a broad ridge with elevations of about 3000 feet during Miocene and early Pliocene time (Axelrod, 1956). By late Pliocene time the range stood at approximately its present elevation (Christensen, 1966).

CASCADE MOUNTAINS

The Cascade Mountains continue the topographic high of the Sierra Nevada northward to the Coast Mountains of British Columbia. The Cascade Mountains are a Cenozoic magmatic arc with a very complex history (see Hammond, this volume). The interbedded marine and non-marine rocks of the Paleogene ancestral Cascade Mountains were folded and faulted during intrusion of early to middle Miocene plutons associated with coeval volcanism. The ancestral Columbia River maintained its course through the mountain chain during this deformation.

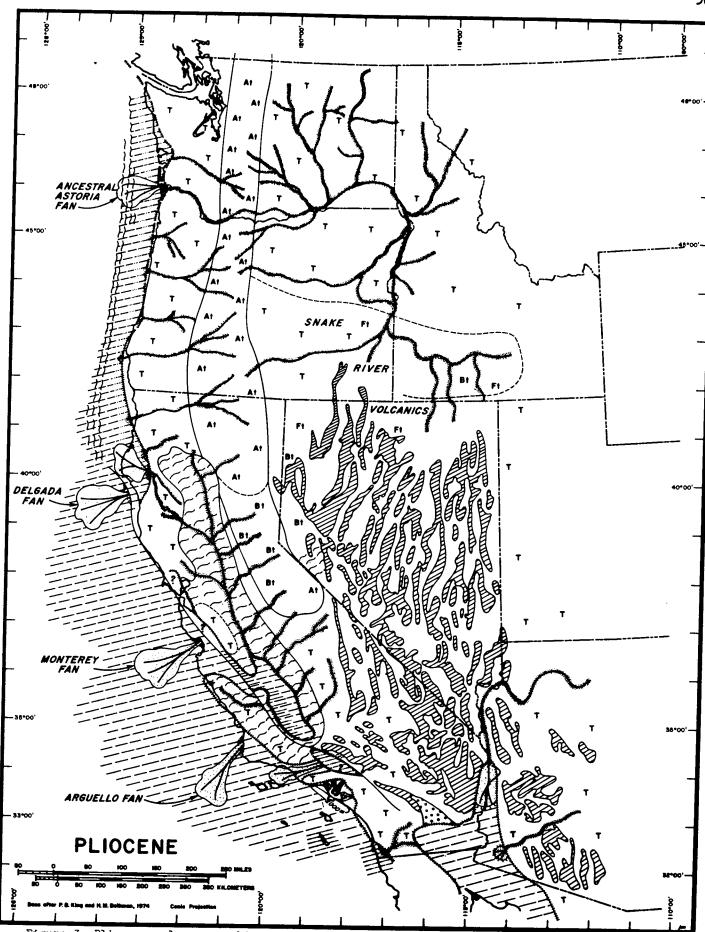


Figure 3. Pliocene paleogeographic map of the western United States. See Figure 2 for symbols.

Middle Miocene Columbia River Basalts flowed from the eastern 'Columbia Plateau westward through the ancestral Columbia River drainage system and into the Puget-Willamette Lowland province. Along the eastern margin of the Willamette Lowland and in the western foothills of the Oregon Cascade Mountains the Columbia River Basalt flows mingled with locally extruded lavas of similar composition (Peck and others, 1964; Snavely and others, 1973).

During the late Miocene and Pliocene Cascade Mountain volcanism continued in response to subduction to the west.

Pleistocene events within the Cascade Mountain province include continued uplift and erosion, formation of the High Cascade stratovolcanoes, and alpine glaciation.

PUGET-WILLAMETTE LOWLAND

West of the magmatic arc of the Cascade Mountains is the Puget-Willamette Lowland province. This topographic low is a structural downwarp between the Cascade Mountains and Coast Range uplifts and extends northward into the Frasier Lowlands of British Columbia. The Puget-Willamette Lowland is topographically analogous to the Great Valley of California although the two lowlands are separated by the Klamath Mountains and have different geologic histories. Outcrops in the Puget-Willamette Lowland province are limited to local uplifts with numerous cross folds (Baldwin, 1947; Snavely and others, 1958; Volkes and others, 1951).

Early Neogene deposits of the Puget-Willamette Lowlands consist of interbedded marginal marine to non-marine volcaniclastic rocks derived from the Cascade Mountains. Locally, middle Miocene basalts were extruded along the eastern flank of the Willamette River Valley (Snavely and Wagner, 1963; Peck, 1964). Basalts of the Columbia River Group flowed westward from the Columbia River Gorge into northwest Oregon and southwest Washington where they were interbedded with shallow water marine sediments and locally derived basalts (Snavely and others, 1973).

Today, the Puget-Willamette Lowland is carpeted with Pleistocene fluvial, lacust-rine, and alluvial sediments, some of which are glacially derived (Easterbrook, this volume).

GREAT VALLEY

The Great Valley, a structural depression between the Sierra Nevada and the Coast Ranges of California, terminates to the north against the Klamath Mountains, and to the south against the Transverse Ranges.

Marine and continental sedimentation occurred in the Great Valley province throughout the Miocene and Pliocene (Addicott, 1968; Huffman, 1972). During the early Miocene the Great Valley province had water depths greater than 6000 feet (2000 meters) with the deepest portion in the southwest corner of the basin adjacent to the San Andreas Fault. Sedimentation rapidly filled the basin to bathyal depths in the middle Miocene and neritic depths in late Miocene and Pliocene (Bandy and Arnal, 1969). Turbidite sedimentation predominated in the center portion of the basin with contemporaneous fluvial and shallow marine deposition to the north, east, and south (Hackel, 1966). The basin extended west across the San Andreas Fault system with a deep water connection to the ocean. Continental sedimentation progressively filled the basin southward and by the Pliocene only a small portion of the basin was still receiving marine sediments. The basin was essentially filled by the end of the Pliocene.

Pleistocene deformation associated with the San Andreas Fault system has uplifted much of the western Great Vally and unroofed some of the Cretaceous outer arc basin sediments of the Franciscan terrain (Hamilton, 1978).

COAST RANGES AND CONTINENTAL SHELVES

The Coast Ranges and continental shelves of western North America consist of accreted marine shelf and slope sediments, intercalated volcanics, and deep sea sediments. The tectonic style of the margin can be used to subdivide the province into three distinct sub-provinces. The subprovinces of the Coast Ranges and continental shelves are 1) the extensional-compressional system of the California Borderlands; 2) the vertically dynamic right lateral strikeslip San Andreas fault system of the California Coast Ranges and continental shelf; and 3) the east-west compressive system of the Oregon and Washington Coast Ranges and shelf.

CALIFORNIA BORDERLANDS

The California Borderlands is the area south and west of the Transverse Ranges of southern California. The Los Angeles and Ventura Basins and numerous offshore basins extending to the shelf-slope break are included in this subprovince.

Prior to the middle Miocene, western California was an area of broad marine onlap with locally prograding non-marine clastic wedges. Tectonic reordering of western North America in the middle Miocene resulted in a general recession of marine waters. Southwest of the Transverse Ranges numerous local structurally controlled deep basins formed adjacent to fault bounded uplifts. Rapid subsidence of these basins is indicated by middle Miocene bathyal and

abyssal deposits overlying early Miocene inner neritic and non-marine deposits. Sediments poured in from the structural highs rapidly filling the eastern basins and spilling over to successively more westward basins (Yerkes and others, 1965; Campbell and Yerkes, 1976).

The tectonic reordering of the California Borderlands has been related to changes in plate interactions (Atwater, 1970; Dickinson, this volume). The style of Borderlands deformation indicates a dominantly extensional stress field associated with basaltic extrusion, uplift of the Transverse Ranges, major left lateral and right lateral slip along faults, and crustal extension giving rise to deep Neogene basins in the Los Angeles, Ventura, and offshore Borderland areas.

Subsequent Pliocene-Pleistocene north-south or northeast-southwest compression has resulted in north-over-south reverse oblique left-lateral faulting at the southern boundary of the Transverse Ranges (Campbell and Yerkes, 1976). The compressional stress system continued into the Pleistocene as evidenced by the warping of the uplifted Pleistocene marine terraces.

CALIFORNIA COAST RANGES AND CONTINENTAL SHELF

North of the Transverse Ranges and the California Borderlands are the Coast Ranges and continental shelf of California. This geologic province is flanked on the east by the Great Valley; to the north, the California Coast Ranges abut the Klamath Mountains.

Neogene and Paleogene rocks of the California Coast Range are superimposed upon two markedly different basement complexes (Hamilton, 1978). West of the San Andreas Fault and north of the Transverse Ranges the basement complex is a granitic-metamorphic terrain referred to as the Salinian Block. East of the San Andreas Fault the Salinian Block may extend as far north as the Mendocino Fracture Zone. East of the San Andreas Fault the basement is a complex terrain of metasedimentary, sedimentary and ophiolitic rocks in places deformed into a melange. This is the Franciscan terrain. Depositional and tectonic models suggest that the Franciscan terrain represents accreted crust and the Slainian Block represents a continental slice transformed northwestward along the San Andreas Fault system (see Dickinson, this volume).

The pre-middle Miocene paleogeography of the California Coast Ranges and continental shelf was dominated by broad marine embayments. Beginning in the middle Miocene the impingement of the East Pacific Rise initiated uplift of the area. The collision of the Rise with the North American continent progressively shifted from a convergence margin to a strike-slip transform margin (Atwater, 1970; Dickinson,

this volume). By the late Miocene the region had subsided and was agin characterized by marine onlap. In the late Miocene and continuing through the Pliocene and Plesistocene, the Coast Range and continental shelf province was again uplifted. Structural basins were filled from the east shifting depocenters further and further west.

Cenozoic shelf sedimentation occurred in structurally controlled basins. Neogene and Quaternary submarine canyons cutting the shelf have allowed by-passing of sediments into deep marine base-of-slope fans (Clark and Greene, this volume).

The strike-slip and vertically active tectonic style of the California Coast Ranges and continental shelf continues today. Pleistocene events are recorded principally by a series of uplifted marine terraces offset by left-lateral faults.

OREGON-WASHINGTON COAST RANGES AND CONTINENTAL SHELF

The Oregon-Washington Coast Ranges extend from the Klamath Mountains on the south northward to Vancouver Island, and are bounded on the east by the structural downwarp of the Puget-Willamette Lowland.

The Cenozoic sedimentation pattern of the Oregon-Washington Coast Ranges was dominated by accretion of marine sediments and intercalated volcanics (Snavely and Wagner, 1963; Snavely and others, 1977). Recent paleomagnetic studies suggest that some of the Paleogene volcanics are accreted "mini plates" (Simpson and Cox, 1978; M.E. Beck, 1979, personal communication; Ray Wells, 1979, personal communication). By the Neogene the "mini plates" had been accreted and the Coast Ranges were being uplifted. Shallow marine embayments penetrated the uplifted Coast Range during the early Miocene at Cape Blanco, Coos Bay and Newport, Oregon, along the mouth of the modern Columbia River, and in the Grays Harbor and Straits of Juan de Fuca areas, Washington.

Although the compressive tectonics system related to sea floor spreading and subduction persisted throughout the Cenozoic in Oregon and Washington, middle Miocene changes in rates and vectors caused major changes in paleogeography (Atwater, 1970; Dickinson, this volume). Uplift of the Coast Ranges became more pronounced. Marine embayments became more restricted and were rapidly filled by prograding non-marine clastic wedges. The early middle Miocene ancestral Columbia River delta and submarine fan were displaced northward as the north plunging end of the Oregon Coast Range anticlinorium shifted northward (Niem, 1973; Niem, 1979, personal communication). Local vents extruded middle Miocene basalts coeval with and geochemically similar to basalts of the Columbia River Basalt Group (Snavely and others, 1973). In the area of the present

lower gorge of the Columbia River basalt flows coming from the Columbia Plateau are interbedded with non-marine and inner neritic sediments (Snavely and others, 1973).

A single subaerial Columbia River Basalt flow extended northward from the Columbia River into central southwestern Washington where it entered the ocean and continuéd westward as a submarine flow (Pease and Hoover, 1957; Snavely and others, 1958; Gower and Pease, 1965; Wagner, 1967).

The Coast Range of southwest Washington was uplifted during the late Miocene into Paleocene cored faulted anticlinal "hills." These "hills" separated shallow marine embayments filling from the east with nonmarine clastic wedges.

The Olympic Mountains, at the northern end of the Coast Range, are a unique subprovience of the Coast Ranges of Oregon and Washington. Principally a late middle Miocene uplift, the Olympic Mountains formed by major underthrusting of shelf and slope sediments accreted as a consequence of sea floor spreading from the Juan de Fuca Rise (Rau, 1973). The Neogene terrain along the western flank of the Olympic Mountains includes allochthomous blocks of both deep marine Neogene rocks and much older rocks. Piercement structures are common along the coast and on the continental shelf west of the Olympic Mountains (Rau, 1973; Snavely and others, 1977).

Superimposed upon the Oregon-Washington continental shelf are late Miocene, Pliocene and Pleistocene structurally controlled depocenters filled with clastic detritus eroded from the uplifted Coast Ranges (Braislin and others, 1971). Late Neogene continental shelf sediments are predominantly siltstone and claystone (Snavely and others, 1977).

Pleistocene modification of the Coast Ranges and continental shelf of Oregon and Washington is principally uplift and emergence as recorded in sequences of uplifted marine terraces. Glaciation within the Coast Range province was restricted to the Olympic Mountains.

SUMMARY

The Neogene paleogeography of the western United States can be summarized as three principle phases:

- 1) An early Miocene continuation of the Paleogene paleogeographic pattern. (see Nilsen & McKee, this volume)
- 2) A middle Miocene tectonic transformation with folding and faulting causing the fragmentation of broader Paleogene and early Neogene paleogeographic patterns.

3) A late Miocene to Recent deformation uplifting mountain ranges and the westward migration of marine depocenters.

The two maps presented in this report are preliminary and present a limited overview on the Neogene paleogeography of the western United States. Future work should include the preparation of maps for time slices representing the early, middle and late Miocene, the Pliocene, and the Pleistocene. Map types should include the distribution of depositional packages of surface accumulated rocks, distribution of faunal and floral provinces, climatic data, and topography and bathymetry.

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